

THE ARCHAEOLOGY OF COASTAL CHANGE, PUERTO RICO

By

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PREFACE

If a world atlas of mythical lands were made, the maps of most centuries would depict at least one lake, legend, bay, or sea with a hidden city in it. Have you would find the watery trail to Atlantis, Lemuria, Mu (which is not short for Lemuria), Camul, Thuleana, Llys Helig, Gair Annabhad, Chao'ia, Lyonesse, the mythical islands of Amerindian legends, the cosmogonic or hidden cities of Galois, and numerous other cities.

Among divers and submariners, legends of hidden cities tend to be reported with considerable conviction. This is not at all surprising, for the unknown is our New World. Not unlike conquistadors, divers are explorers who often mistake real wonders for assumed, archetypal fancies. To divers, the unknown is almost a planet within a planet, a world without gravity, without horizon, without sun, without air. In this liquid realm, the natural and the human-made often blend in perfection.

A sandstone reef may be taken for the wall of a hidden city, particularly in poor visibility. In water over 100 feet deep, the diver's judgment and perception are further impaired by nitrogen narcosis, an euphoria that increases with depth and closely resembles alcohol intoxication.

At King's Bay, Georgia, aboard a USN nuclear submarine open to civilians for the day, the casual remark of my research led an officer to mention "an incredible archaeological discovery in deep water." He

wouldn't go any farther, as the information was supposedly classified "You found a pyrexia," I said. Taken aback, the naval officer probably inquired about my source of information. I said I had read about it somewhere. To this day, I still wonder what his reaction might have been, but I told him about the power of myths, a power far greater than the nuclear warheads aboard his Grey Lady.

There is but one of various instances in which narrow, vertical margins and doors have told me of "mysterious races" in the sands of the sea. In it all history, no more Spanish archeologists claim what coincided with the possibility that Tartessos (the Biblical Tarshish?) was an actual city link to the sea? Or is there, perhaps, a geoarchaeological issue behind the legend? The draining of Galton's Azule Lagoon revealed no signs of the fabled city that was supposed to be there (Ginspre 1948: 60). On the other hand, the legendary wooden town in The Philippines Lake Taal was recently proven to be true (Ginspre 1966).

The theory of plate tectonics indicates that wooden continents are a geophysical impossibility. This is due to the lightweight composition of continents in relation to the denser crustal floor. From a contemporary geological perspective, the idea of a wooden continent is as *Ma or Atlantis* is as absurd as a wooden subway, or a wooden ice cube in your favorite drink. Just as the subway floats on air, the continental masses "float" over the denser rock beneath them (Roberts 1977: 24).

Nonetheless, wooden masses are a fact. Wooden islands are a fact. Crosses in the sea are a fact. As for the Deluge, which is found in religious traditions throughout the world, an oceanographer of the nature of Devere

Englund has suggested it may refer to the Helocene marine transgression (Englund et al. 1973), which flooded the world's seashores after the Wisconsin glacial era. That is a challenging view that many archaeologists would reject in favor of localized floods, combined with diffusion of legends (Garrow 1972: 384).

The problem with Englund's hypothesis is not the reality of the Helocene transgression, amply proven by earth scientists, but rather assuming to mythic an event that began some 12,000 years ago and which lasted thousands of years. As a witness to a recent hurricane, I can well understand local legends in prehistory may have been strong enough to create apocalyptic myths.

There again, in favor of Englund's hypothesis, there may well have been a stage in the Helocene transgression fast enough to become a 'cultural event' (Pinnock 1972: 27). If that is the case, then the Biblical Deluge is essentially true, which is not to say that only Noah's family survived it. As scientists, we should recognize the great amount of symbols and historic knowledge that is contained in religious writings regardless of how poetic the presentation.

Perhaps the mythical and scientific views on ancient legends have never come closer than at Bimini Road, an underwater pavement off Bimini Island in the Bahamas. The story goes back to 1968, when psychic leader Edgar Cayce predicted that, somewhere between 1800 and 1900, parts of Atlantis would be rediscovered near Bimini. And indeed, in 1968, the new famous "road" was discovered (Sherkin 1974). Other expeditions followed, revealing additional walls and pavements. These underwater discoveries thrilled the Atlantologists, and called the attention

of prehistoric cartographers and cartographers alike, such as Deniro, Schiavelli, Jacques Cartier, and Alex Landberg (1976) of the *In Search Of Unknown* series.

But exactly what is the nature of the *Shimau Complex*? To begin with, the famous natural "road" is bedrock, naturally formed at the structural zone (crust + lower sea level). The "unpolished blocks" of the "road" are boulders, but they suspiciously resemble a terraced pavement smoothly cut by natural water weathering. At least some of the random walls in Bahamian waters appear to be the marks of historic turtle breaks, and the "wooden temple pillars" look very much like speleothems from submerged karst caves. As for the artifacts supposedly recovered by divers, I have yet to see the photograph of an *Africanus* pottery shard, or any prehistoric pottery shard for that matter, found at Shimau Road. To this day, there is no conclusive evidence that a prehistoric civilization was destroyed by a massive catastrophe in the Bahamas. Then again, there is no conclusive evidence against it.

Last I be taken for an anti mystic, I should say that my own interest in mystic lands originated as a childhood vision. The fantastic daydream led me to Poe, Byron and Yeats, and further intensified my passion for the unknown. Strangely enough, my favorite real garden was the same roof where I would later discover a mystic prehistoric hamlet (Vega 1971, 1982). As a scientist, it took me years of great intellectual and physical effort, and the retrieval of hundreds of artifacts and evidence in the process, before I was able to ascertain that I had found the actual ruins of a mystic prehistoric village—as opposed to the secondary deposition of artifacts washed from the nearby coast, or a Bermuda Triangle mirage.

My discovery was a far cry from Poma City in the Sun. Instead of "temples and palaces and towers," I found pottery shards, stone axes, and a small carved face. But that rather plain misshapen shell within the first to be excavated in the Caribbean—showed me a new way of looking at prehistory.

Continents are not fixed boundaries. They are landforms in eternal change, drifting back and forth across island and continental shelves. In the final analysis, the myth of unbroken lands merely a fundamental truth: what was once dry land may be sea-bed now.

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Archaeological interpretation of testate and ostracite land-use level changes indicates a highly dynamic Holocene environmental history for the Caribbean Archipelago. By 15,000 B.P., sea level may have stood 100,000 m below present mean sea level (MSL). Benthometric and paleontological data show that the Caribbean offshore islands were not attached to the American mainland during or after the Wisconsin glacial maxima. A lower sea level facilitated a maritime entry of early prehistoric peoples into the Caribbean. It is hypothesized that the first migrants were maritime hunter-gatherers who sustained offshore in response to the gradual denudation of the continental megashore. The testate component of relative sea level is estimated for the main island of Puerto Rico as a case study. Geoarchaeological interpretation of 25 prehistoric and 3 historic littoral sites indicates a MSL downward tilt (a) and since prehistoric times, as defined from the dated positions of prehistoric middens on the south coast, and the absence of prehistoric sites on the north coast, except for the prograded outlet of the Luquillo River. The roughly uniform, estimated shoreline migration (B&B) of ancient sites on all four islands suggests that the testate tilt named by B&B

B.P. High marine transgressions of Delwood corals sites and Native military installations suggests a global sea level rise of 1 m or more since 1600 B.P. Preliminary geoarchaeological research for the Virgin Islands supports the historic unity of these islands and Puerto Rico. Submerged prehistoric sites are predicted for the north coast of Puerto Rico, Virgin Gorda, Anegada, and possibly the other islands in between. In these areas, prehistoric mounds averaging 25 m in diameter should be found on the surface vicinity of beachrock pavements, at a water depth of 5 m or more.

CHAPTER 1

THEORETICAL FOUNDATIONS

Introduction

These things are far more complicated than they are, and not to be wondered at if in our days the raised coasts of these islands occurred in many countries, and if moreover such coasts ever existed. The seas, the continents, the bridges of the waters, the changes of language told of lost, time-estimated, every vestige of the past.

Camden 20 (1602): 1432, 1544 (1606: 37)

Use the sea level, the shore facilities on the coast of L, the place of our first ancient knowledge.

Richard Coeur (1895: vi)

As early as 210,000-400,000 B.P., Lower Pleistocene humans explored the edge of the sea in search of shellfish at Terra Amata, near Southern France (de Benkey 1980; Wile 1983). Fully modern humans appear in the archaeological record by 35,000-40,000 B.P., but it is only by c.9000 B.P. that the first coastal cultures appear (Cohen 1987: 131). Today, there is growing recognition that this may well be a biased picture created by the subsupercession of coastal sites (Pienkowski 1985; Suppl 1986a: 30). Perhaps there were shellfish eaters at the same time, or before, there were big-game hunters, but their sites are now under water. In the last 77,000 years, sea level may have risen an astounding 100-130 meters (330-430 ft.), drowning millions of square miles of the world's coastal zones (Kennedy 1981: 264; Harman et al. 1979: 265; McManus and Smey 1985: 1121).

Myths and folklore tell us of towns and kingdoms drowned by the sea. Some of these stories are pure fancy, others are based on actual events, such as the dramatic sinking of the "ghost city" of Port Royal, Jamaica, by an earthquake in 1692 (Stanley and Woodward 1984).

Since the 1850s, archaeologists speculated that prehistoric man lived on portions of the now submerged continental shelf (Ellis 1937). It was assumed, however, that such sites would have been destroyed by the rising seas of the Holocene transgression. Today, the archaeological community is beginning to recognize that the traditional "land under hypothesis" is false. In the last two decades, diving researchers have proven that ancient, land sites of virtually any age may survive inundation, in both fresh and salt water environments, and be systematically studied under water.

Submerged land sites of prehistoric, classical, or historic cultures have already been found in Florida (Coker 1969, Suppe 1980a), California (Masters 1985, Morarty 1981), Maine (Richardson 1985), Jamaica (Mayer and Meyer 1972), Puerto Rico (Sage 1961, 1962), Nova and St. Bartolom (Mey 1973), Denmark (Anderson 1988), Russia and Yugoslavia (Radt 1988), Sweden (Larsson 1988), France (Gallin, et al. 1983, Trugnot, et al. 1982), England (Ellis 1984), Spain (Larsburg 1979, Mars 1988), Portugal (Claudio Bontlele, personal communication 1988), Gibraltar (Woodward 1984), Italy (Lowe 1978), Greece (Gifford 1988, Fournel 1987), Sardinia (1988), Turkey (Flannery, et al. 1973, Garber 1979), Tunisia and Egypt (Martin and Flannery 1977), Libya (Flannery 1979, 1983), Israel (Ellis 1983), Lebanon (Frost 1978), Tanzania (Jones 1984), The Philippines (Rogerson 1988), etc.

Research Goals

We started discussing coastal change centuries before the change to land right off the shore
Linda Quidley (personal communication)

The brown Indians and the gardens of the sea, months later and the work they were
all building just we were that one thing too
John Barlow (c. 1940-1970)

The original idea behind this study was to search for submerged and intertidal archaeological sites in Puerto Rico. Departing from the author's previous discovery and excavation of a submerged prehistoric site off Puerto Rico's north shore (Maga 1981), it came as a natural progression to search for other sites, to demonstrate that the immersion of Caribbean archaeological sites is the rule, not the exception.

As other submerged and intertidal sites were found and analyzed, a formal theory of Caribbean coastal change began to take shape. Reexamination of the literature formally revealed what the author had previously suspected from observation: that Caribbean archaeology has wrongly presented the coastal environment as a passive actor, as an unchanging landscape that yields shellfish and sponges.

Studies of Caribbean prehistoric migrations are characterized by only cursory references, if any, to Holocene sea levels (Dorán 1988, Cronin and Brown 1988, Price 1971, Rapp 1978, Brown, 1984, 1986, Brown and Allmon 1979). Among the few exceptions are the pioneer papers by Melikian (1976a, 1976b), and Rapp (1982b), dominated by whom.

Beyond sea level per se, the region has produced few studies linking archaeology to geomorphology, clearly validating Webster's (1983) criticism that Caribbean archaeology is characterized by a significant territorial bias.

While changes in land ecology are deemed central to prehistory, changes in marine ecology remain largely unexplored.

Extensive personal observations confirm that most native Caribbean archaeologists are not particularly situated to or knowledgeable of the sea (Watters 1993: 522), an ironic situation worthy of archaeological inquiry. The irony is not simply that we are islanders, but more importantly, that the prehistoric peoples whose lifeways we are trying to reconstruct lived in close physical and spiritual bond with the sea.

With these problems in mind, three research objectives were envisioned: to develop a maritime perspective on Caribbean prehistoric megafauna (Chapter II), to search for submerged sites in Puerto Rico as a case study (Chapter III), and to integrate the field data into a regional model of Caribbean coastal geomorphology (Chapter IV).

Defendants of the Prison

It is said that within a few years Puerto Island has been reduced from being called as bright as day, and that of three lighthouses built on it none is left. Not have been washed away and the land will surely erode.

OSCAR CASTELL BLANCO, 1900 (1984: 26)

I have heard it said that some people do not believe in the existence of slaves in the sea. What a pity not to believe in those which are no stronger in any legend.

HERNÁN G. FLORES (c. 1975, n.d.)

Submerged land sites present a number of unique problems to the researcher. Where do you dive for them? What environmental and site modifications occurred as a result of emergence? Are the artifacts on the sea floor an actual site, or the results of secondary deposition due to a site eroded on the nearby coast? How do you go about surveying and excavating such sites? How do submerged land sites help us recognize coastal change?

Lebanese Submerged Land Sites

The first submerged local sites to be studied were Mediterranean classical and pre-classical ports. In the nineteenth century, engineer Giuseppe de Fazio hired sponge divers to survey the ancient harbors of Pegasus and Marone, in the Bay of Naples. In 1806, archaeologist Robert Gauthier continued de Fazio's survey through a glass-bottom boat (Bucki 1988: 134). In 1819, employing hired-hat divers and free fishermen, Gustav Jordaet discovered the ancient remains of the once spectacular harbor of Phoenice near Alexandria, Egypt (Fleming 1975: 40).

That the Mediterranean is the realm of submerged local archaeology is not at all surprising. Fortified villages date back to 8,000 B.P., and the coast is crowded with ruins ranging from Neolithic times to the Roman Empire. Classical ports alone run in the hundreds, many of which have been submerged by tectonic movements in a highly active seismic region (Fleming 1975: 164). Numerous ancient sites are visible from shore, and particularly from the air. In 1924, in a reconnaissance flight along the Lebanese coast, Jewish priest and archaeologist Aristeus Finkelard discovered the ancient ruins of the Phoenician city of Tyre, which he promptly surveyed with sponge divers (Bucki 1988: 135).

Following the destruction of the Convent-Dagupan aqueduct in 1943, scuba diving became a popular Mediterranean sport after the end of World War II. This led to the discovery and exploration of numerous ancient ports. By the early 1970s, divers and archaeologists had located over one hundred "sites in the sea" (Marston and Fleming 1977: 222).

The end of World War II also saw the rise of deep-sea microscopy, producing a quantum leap in sea level studies (Emmons 1975). Combined

with the availability of lightweight scuba gear, and the discovery of Mediterranean submerged land sites, the new oceanographic discipline presented a challenge to prehistoric archaeologists.

Searching for drowned prehistoric sites was first proposed in the United States by Gagan (1963), with early discussions by Schuch (1961), Shepard (1964), Emery and Edwards (1966), Salwan (1967), Bodin (1968) and Warren (1968). At the time, the archaeological community was not ready for diving and sea level research, and the problem remained more theoretical than practical.

In the Caribbean, the search for prehistoric marine sites was first proposed by Nashelen (1976a), following a survey of Antigua, which indicated that shell middens on the NE coast are underpinning marine terraces, while middens on the SW coast are now inland. More recently, Ruppel (1990b) has also encouraged a search for drowned Caribbean sites, based on his own successful underwater research in West Florida (Lightfoot and Ruppel 1986, Ruppel 1989a).

Environmental Variables

Submerged land sites may occur in a wide variety of environmental settings, including marshes (Larsen et al. 1978), lake (Hargrave 1968, Ford 1978), river (Palmer et al. 1981), submerged land areas (Pinnington 1983, 1985, see areas (Woodward 1960), man (Larsen 1982), estuaries (Pruett et al. 1988), semi-protected tropical coasts (Step 1981), sheltered littoral coasts (Clifford 1983), and high energy coasts (Masters 1983, Blache 1990).

Considering these widely divergent environments, one can only agree with Flannery (1983a: 188) that "no single factor in the geomorphology of the coastal zone is necessary or sufficient to ensure preservation of archaeological remains under water."

Three important variables must be considered here: 1) that the modern coastal environment need not be the same as when the site was created, possibly being exposed now but protected then, 2) that the speed of the extreme transgression may be a coastal factor in some instances, with fast flooding increasing the chance of site preservation, and 3) that previous substantial deposition may act as a shield during transgression, even in high-energy coasts.

In a comparative analysis of the preservation of wooden sites world-wide, Flannery (1983a: 180) has concluded that most coastal archaeological sites were originally located in sheltered places formed by the original shoreline, and that most of these sites may survive at least one marine transgression. Undoubtedly, some of the artifacts occasionally found by beach comers are the result of secondary deposition by longshore and rip currents (Shaw 1983: 399). In some cases, artifacts may have fallen from wrecks, or be a part of the weather pattern of a shipwreck site. Nevertheless, there is ample, world-wide evidence that many land sites do survive the process of inundation with little or no damage at all. In the long run, the sea may actually be a preservational agent, providing anoxic conditions in some instances, and protecting the site from the anthropogenic disturbances or deterioration of many land sites: plowing, construction, road intrusions, and looting.

Method and Theory

Underwater surveying and excavation techniques are discussed in detail by Bass (1970), Nute (1984), Polarov (1973), and St. John Wilson (1978). At the macro- and mega-level, excavating a sunken land site is somewhat akin to shipwreck archaeology. There are, of course, fundamental differences in the nature of both sites. In principle, shipwrecks are not stratigraphic sites. On occasion, however, two or more shipwreck sites may be superimposed, or be found within the larger stratigraphic context of a harbor.

For a general overview of shipwreck archaeology, see Bass (1978, 1988), Johnsson (1974) and Thorsenmorton (1987). Theoretical issues of shipwreck archaeology are discussed by Blackwater (1981), Gleditsen (1984), Gosh (1988), Mackelvey (1978) and Murphy (1981). For a general history of underwater archaeology, including both shipwrecks and sunken land sites, see Hoffmann (1987), Mackelvey (1983), and Bass (1988).

In the most elaborate theoretical discussion of shipwreck sites in print, Mackelvey (1978: 190) argues against the traditional 'good-bad' shipwreck dichotomy, indicating that most wreck sites are neither totally coherent nor totally scattered. The same principle is applicable to the stratigraphy of submerged land sites, which may range from intact to completely disturbed. As mentioned elsewhere, the fundamental point is that submerged land sites are not necessarily without stratigraphy (Anderson 1980, Goddard 1988, Flemming 1988a).

At Isla Verde, a submerged prehistoric site off Puerto Rico, the author had no difficulty recovering perfectly vertical pole in situ, using a home-made water dredge (Naga 1981). However, it is an altogether different

story in sand or silt, both of which have a stability angle generally under 45 degrees. As one solution to this problem, Lewis (1974: 147) has used a vertical caisson, large enough to accommodate a diver with a water dredge or hoist for controlled excavations. Another solution has been offered by Gifford (1983), consisting of a down-bell diver. Both techniques have proven quite useful to shore-dwelling researchers, but neither has been tested on densely-packed, underwater shell mounds.

Palaeogeography

One of the fundamental postulates of this dissertation is that archaeological sites may be used as markers of coastal change. This is by no means a new idea, pioneered for Puerto Rico by Kaye (1936). More recently, Fairbridge (1976) has traced Holocene sea levels in Brazil by radiocarbon dating of shell mounds on coastal beach ridges. From an environmental perspective, the study of coastal archaeological sites is relevant to geomorphologists, geographers, and, indirectly, to engineers and planners engaged in coastal erosion and sedimentation control (Berth and Trues 1984, Brooks et al. 1978, Chardon 1977). As advocated by the Commission on the Coastal Environment, understanding coastal change requires "historical sources as well as geomorphological methods" (Ford 1985: 4).

Reconstructing shoreline migration on the basis of archaeological data is an archaeological endeavor, whether it is conducted by an archaeologist, geomorphologist, or geographer. It takes archaeological thinking and technique to define the relationship of structures or features to the ancient shoreline. Sketching a beachmarker for a quay or break wall would produce a completely erroneous evaluation of shoreline migration for

that particular coastline. Indeed, many of the purported sites of Atlantis are nothing more than the erroneous interpretation of sea walls as the remains of ancient houses and temples.

Research Domains

My eyes have faithfully looked down
On the long right line of that house,
That light from out the level sea
Glimmers up the turret slowly
(*Egypt: Atlas Press, 1900-1901*)

Human settlements require terraces for expansion
as is natural everywhere in Cefalonia
(*Islands: Goussakis (2004), 98*)

Evidence of sea level and shoreline migration may be profitably grouped into three research domains: I) geomorphological, II) cultural, and III) historical. These domains comprise particular sources of evidence and have frontiers, each encompassing a shorter but more finely dated sequence of time.

For purposes of this project, the geomorphological domain comprises the Holocene period, roughly 10,000 B.P. to the present. Potential geomorphological indicators of coastal change in the Caribbean include beachrock, sediments, shallow water mollusks, coral, rock ridges, ridges and ridges rock, salt marsh peat, drowned river beds, epibenthos, fossil beach ridges, sea caves, drowned forests, tidal terraces, and sea level rise.

The cultural domain comprises the range of human occupation of the study area, possibly extending back to 7000 B.P. for the Caribbean (Wegert et al. 1988, (Grove 1988) Maggioni and Kruse 1979: 645, Pons 1971, Brown and Allaire 1975: 465, Wilson and Orsney 1979, Wiley 1979). Cultural

indicators of coastal change include prehistoric and historic shell middens, human burials, petroglyphs, pictographs, and isolated artifacts embedded in beachrocks.

The historic domain ranges from the foundation of the first European settlements to the present. For the Caribbean, this period begins in the last decade of the 15th century, with the settlements of Navidad, Isabela, Concepción de la Vega and Santa Domingo in Hispaniola (Kleinman 1990, MacAlister 1994, Parry 1993, Sauer 1986). For Puerto Rico, the historic period begins in 1493, with the foundation of Ponce on the west coast of the island, by Juan Ponce de León, followed in 1508-9 by the settlements of Caparra and Guayan, on the north and south coasts respectively (The 1994: 34).

Historic markers of coastal change include port installations, coastal fortifications and fortification sites, historic and modern maps, literary descriptions, railroad tracks, aerial photographs dating back to the 1930s, historic artifacts embedded in beachrocks, and modern coastal structures in general.

CHAPTER 2

CHANGING LEVELS OF LAND AND SEA

Marine transgressions may occur due to a variety of local, regional, or global processes. Land may be created or subside due to tectonic movement, postglacial rebound, isostatic depression, and sediment compaction (Brace 1971, Mercer 1971, 1980, Murray 1982). Sea level may rise due to changes in the volume of ocean water or of ocean ridges (Dowson and Jones 1976, Clark et al. 1978, Coleman and Smith 1984, Curry 1981, Rindum et al. 1979). Following current theory, the major agent of sea-level change has been a global or eustatic sea-level rise due to the melting of land-blocked ice, possibly resulting from variations in the earth's orbit (Haye et al. 1978: 1121).

Sea level remains a controversial topic of vast complexity. An impressive body of literature has been accumulated (Clark and Fairbridge 1982), but there is serious disagreement on the chronology and boundaries of Quaternary sea levels. When did the sea reach its present level? Did it rise steadily, or in an erratic fashion? Did it rise above present mean sea level during the Holocene transgression? These questions need to be addressed at the regional level, as global sea level curves are not sensitive enough to account for local and regional variations (Clarkson et al. 1978, Chappell 1974, Chappell and Thom 1977, Coles 1978, Fairbridge 1979, Murray et al. 1976).

Glacial and Tectonic Eustasy

In the six hundredth year of Nineteen Six, on the six hundredth day of the second-moment-in that day at the springs of the great deep, hand built, and the fountains of the fountains were opened, and man fell on the earth body days and body nights.

The Bible, Genesis 6:11

The gods feared the flood, they fled. They climbed into the forests of Asia. The gods remained like a dog on the wall. For six days and nights water and flood continued on the horizon, inundated the land. And at midnight rain fell on that

Epic of Gilgamesh (in Cowley 1972:21-2)

Eustatic movements of sea level are produced by changes in the volume of ocean water, or changes in the volume of ocean basins. In theory numerous factors may alter the volume of the oceans, including the production of juvenile water and its reabsorption in the earth's mantle, changes in the moisture content of the atmosphere, desiccation of ocean basins, changes in mean oceanic temperature, and changes in the volume of land-locked ice (glacio-eustasy). There is a general consensus that glacio-eustasy is by far the most significant of these processes, the other factors having either no or an impact on Quaternary sea levels as far as the problem is presently understood (Dowson and Jones 1979: 187).

In the case of mean water temperature, a rise of 1°C would warm sea level by a mere 0.08 m (West 1979: 362). Deep-sea core studies indicate that six million years ago, the Mediterranean became isolated from the rest of the ocean, dried up, and turned into a desert (Van 1972: 27). This may have resulted in a significant rise in sea level of some 12 m (COP 8:1). However, there is no evidence that a desiccation of that magnitude occurred during the Quaternary Period.

During the last two million years, the earth has undergone some twenty ice ages, in cycles of roughly 100,000 years (Flannery, 1982: 18). These glaciations are caused by a fall in atmospheric temperature, resulting

an increased snowfall, reduced ablation (glacial retreat), and increased freezing of ground water (Weert 1973: 135). With each glaciation, water is abstracted from the ocean, resulting in a drop in sea level (Figure 1). The process does not include the formation of icebergs and ice shelves, as these floating bodies of ice displace their own mass of water.

The cycles of cold (glacial) and warm (interglacial) periods is one of the fundamental mysteries of the earth. Proposed explanations include variations in the output of the sun, partial blockage of solar energy by interstellar or volcanic dust, changes in the deep circulation of the oceans, continental drift from arctic poles to tropical regions, etc. Current theory suggests that these climatic variations are caused by changes in the earth's orbital geometry, which occur in thousand cycles of 100,000 years (Hays et al. 1976: 1121).

Significant variation fluctuations in sea level may also be caused by changes in the volume of ocean basins. The primary factor of basin volume is the growth of mid-ocean ridges, which form an underwater mountain chain that runs for more than 80,000 kilometers (50,000 miles), and rise an average of 3,000 to 4,500 ft.) above the adjacent plain (Robert 1977). Vertical movements of ocean ridges may result from sea floor spreading, producing a gradual sea level rise of 1 cm (1/32 in.) per 1,000 years (Dewar and Jones 1976: 196).

In principle, the capacity of ocean basins may also be reduced by the deposition of land-derived sediment. The Mississippi River alone discharges over 500 million metric tons of sediment per year (Everts 1973: 48). However, much of this effect is cancelled out by isostatic depression of crust beneath the added load, which may cause localized subsurgences of

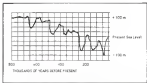


Figure 1. Glacial maximum sea level change for the past 80,000 years (after Mörhacik 1978, p. 21)

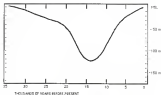


Figure 2. Sea level curves for the U.S. Atlantic Continental Shelf for the last 20,000 years (after Milliman and Meadey 1983: 1182, Fig. 1)

coastal areas through downwarping, as discussed in the following section. Finally, changes in the shape of the earth may introduce considerable regional differentiation into the eustatic sea level curve (Kamnet, 1968-1969). This global factor is controlled by the earth's rotation and mass distribution, and in theory, may cause a rise in sea level in some regions and a drop in others (Dawson and Jones 1979-1981, Mörner 1984).

Barriers and Local Processes

Atlantic was an island larger than Libya and Asia put together. Though it was consequently overwhelmed by catastrophe and to the horror of the superstitious mind which prevents the free passage of time and ridges of the world into the open sea.
 Paus. 1.407-442 9-15 (1971-198)

The myth of the Deluge, Noah, Nostradamus, Deluge and Niagara refer to an event in the Atlantic Ocean "beneath the surface" which sank the Atlantic way back into of what had been there.

Julius Rosenberg (1991,199)

Marine transgressions may occur at specific instants due to land movements independent of eustatic sea level. Blocks of land may subside along geological faults due to slight tectonic changes. Movement may be gradual or occur in a sudden, seemingly apocalyptic landslide, such as the catastrophic destruction of Haida in Greece (Mörner 1990-1991), or Janusova in the Caribbean island of Haiti (Marr 1992-1993). Johnson (1994) has compiled a listing of eastern Caribbean earthquakes since 1500. Catastrophic landslides may also be caused by volcanic eruptions, as the earth's crust readjusts to the rapid outpour of lava (Mörner and Fleming 1977-1981, Butler 1984, 199).

Land movements may also result from the loading or unloading of ice, sediments, or water. Parts of Scotland, Scandinavia, and Canada, once glaciated, are currently being uplifted as a response to ice unloading, with a

corresponding downward movement in other areas (West 1977: 151). Deltaic areas may be gradually submerged by constant depression under the heavy load of river sediments (Coleman and Smith 1944). This process may be further compounded by sediment compaction (Pienning 1977: 209).

Storm surges are capable of producing localized marine transgressions of devastating power. In 1880, at Galveston, Texas, a hurricane surge raised the sea to over 4.5 m (15 ft) above its normal tidal range, on top of which giant waves invaded the city (Stanton 1966). Williams (1980) and Salvia (1972) have compiled chronological listings of Caribbean storms and hurricanes since 1492.

Significant coastal modification may result from erosion, with or without changes in the relative height of land and sea. Hollersom, Enderwh, Winkleson and numerous other English settlements are now under water due to erosion (Morrison 1989: 132). A similar fate happened to Tyndale, on the north coast of Sully (Pienning 1972: 134). Anthropogenic factors may accelerate the process of sediment loss. The construction of breakwaters, quays, etc., the extraction of beach sand, and the clearing of the coastal flora may all result in massive, localized erosion (Charlton 1977). Contrary to the opinion of many archaeologists and geomorphologists, eroded sites may reverse the process of immersion. In Dorset, England, drains working in one locality have uncovered parts of a medieval church, off an eroded coast (Stanton 1966: 21).

Latin/Queniaman Sea Level History

Deluge: The great flood that appears in the religious traditions of virtually all peoples and that destroyed all but a few remnants of the earth.

Charles Henry (1975: 142)

to select those when/where was/led on the planet, a great first step to theory of
 that which was not the story (so that the world could be rebuilt)
 Malvern, Michael (London, in British 1929,200)

Sea level curves have been produced by Blackwelder et al. (1978),
 Bloom (1967), Clark et al. (1974), Curry (1981), Fairbridge (1981), Kelson
 (1942), Milliman and Meade (1983), Morner (1971), Schell and Stuever
 (1987), Shepard (1983), and others. On the basis of the terms, models, and
 constant variables discussed above, it is clear that an exact, global sea level
 curve is impossible. Thus, the concept of unity should be understood as an
 average sea level (Walcott 1971: 16). For the archaeologist, the problem
 implies difficulty in matching sea level curves with cultural chronologies
 (Shepard 1983a: 42). In other words, it is impossible to have a strict, global
 correlation between depth of submerged land sites and their ages (Table 1).
 There is general consensus that the sea was near its present level by about
 55,000 B.P. (Figure 2). Then it began to drop as the Laurentide and other
 Pleistocene ice sheets expanded over the Northern Hemisphere (Bardine et
 al. 1975: 1082). By 35,000 to 15,000 B.P., sea level reached its lowest point
 in the Wisconsin glaciation. Estimates for this lowest sea level stand
 include minus 60 m (Blackwelder et al. 1978), minus 85-90 m (Morner
 1971), and minus 100 m (Milliman and Meade 1983). By 17,000-15,000
 B.P., the Holocene transgression began, with a rapid rise in sea level at
 the order of 4 mm (4 in.) per year, until about 7,000 B.P. (Kornath, 1942:
 276). By that time, sea level may have been more 10 m (32 ft.) below present
 level (Shepard 1983).

From 7,000 B.P. onwards, the sea continued to rise at a much slower
 pace. At this point, there are important disagreements among researchers

TABLE I
Global Correlation of Prehistoric Submerged Land Sites

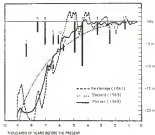
	Site Location	Depth (Meters)	Cultural Affiliation	Years B.P.
1	Forest of Caminal (Denmark) and Sweden	0-8	Mesolithic	6000-4000
2	Norwoudy A (NW France)	0-7	Neolithic	4000-3000
3	Norwoudy B (NW France)	0-8	Iron Age	4000-3000
4	Norwoudy C (NW France)	0-8	Iron Age (Salt Works)	3000-2000
5	Lensvik (South of France)	0-9	Neolithic	1800-4000
6	Kellinick Bay (Spain, Greece)	10	Neolithic	1800-4000
7	Syros Psephos (Syrothian, Greece)	0-10	Neolithic to Iron Age	1800-3000
8	Island A (Mediterranean Coast)	0-4	Neolithic to Chalcolithic	4000-3000
9	Island B (Mediterranean Coast)	0-8	Neolithic	4000-3000
10	Enlène Vallée (U.S.A.)	0-8	La Jolla (Prehistoric)	4000-4000
11	West Florida A (U.S.A.)		Pinckney Island (Cemetary)	1000
12	West Florida B (U.S.A.)	0-8	Mobile Indian (Prehistoric)	1000
13	North Coast, Puerto Rico (Caribbean)	0-10	Orinoco (Cemetary)	1000-400

(Source: Masters and Fleming 1963: 697, Suppl. 1966a, Vaga 1944.)

(Figure 3). Van Andel and Lohr (1964), Fairbridge (1961), and others have proposed sea level fluctuations with higher-than-present sea levels in the late Holocene, while Curry (1961), Schell and Stover (1968), Shepard (1963) and others have reported evidence of a steady sea level rise approaching the present level by 4-800 to 3,000 B.P. Modern corroboration from recent strategic analyses of glacial history support slightly higher late Holocene sea levels than at present (Kamster, 1982: 275), but the problem may well remain an open question.

The rising seas were marked by interruptions or stillstands. These periods of stationary sea level allowed the development of marine terraces, which may be observed under water. Prominent submarine terraces are found in the Caribbean, the Gulf of Mexico, and the Pacific at an average depth of 15 m (49 ft.) and 40-60 m (130-195 ft.). Of course, the depths of marine terraces may be affected by local isostasy, so that Pleistocene terraces may be found in shallow water and even above present sea level (Osney 1960).

Stillstands also allowed the formation of sea caves, beaches, and associated geomorphological features. The archaeological importance of stillstands cannot be overestimated, for these periods of stationary sea level made possible the formation of permanent or semi-permanent, windified coastal sites, as discussed in Chapter 3.



Black bars represent the depth range and median of submergent prehistoric sites as outlined in Table 1.

Figure 5. A comparison of three constant sea level curves, for the last 10,000 years—after Fairbridge 1974: 322, Fig. 1.

CHAPTER 3 CARIBBEAN PREHISTORIC MIGRATIONS

The Sea and the Islands

They sailed because they were afraid or careless, because they were hungry or
curiously. Because they fell San Felipe to sea and lost San Felipe. There was a
rescue for each man — they were coming to find something or leave something, or
get something to dig up something or bury something or leave something alone.
Ray Bradbury (1951: 12)

I must remember to use substandard verbs.
Arvid Moberg Lindbergh (1976: 122)

The Caribbean Sea is separated from the Atlantic Ocean by a chain of islands known as the West Indies. Political and historic divisions aside, the Caribbean Archipelago is broadly divided into five groups: 1) Bahamas-Turks-Caicos Islands, including Bonaire, Aruba, and Curacao; 2) Greater Antilles: Virgin Islands, extending east to Anguilla Passage; 3) the Lesser Antilles, from Barbados to Grenada, and including the Dutch islands of Aruba, Curacao, and Bonaire; 4) the Continental Islands, including Trinidad, Tobago, and the Marguerite Island group; and 5) the Grand Cayman-Banco Verde group, including Matamoros and Banco Verde.

Structurally, most of the region lies atop the Caribbean lithospheric plate, which interacts with the Cocos plate to the east, and the much larger North American and South American plates (Figure 4). For a general discussion of plate tectonics, see Brillot (1994), and Roberts (1993). The

evolution, and movement of the Caribbean plate are discussed by Mollet and Dinkelman (1972), Sykes et al. (1982), and Parker et al. (1986).

The northeastern boundary between the Caribbean and North American plate is marked by the Puerto Rico Trench. Nearly five miles NE of Puerto Rico, the Trench reaches a depth of 9,319 m (30,545 ft), the deepest zone of the Atlantic Ocean. To give an idea of its depth, the Puerto Rico Trench has enough vertical space to accommodate the 8,848 m (29,030 ft) of Mt. Everest (Sediment Cores 1979: 2). This is a seismic zone, characterized by a large concentration of earthquakes but no volcanic activity since the Oligocene (Parker et al. 1986: 138). In contrast, the southern portion of the island arc has exhibited volcanic activity as recently as 1871, at St. Vincent's Soufriere (Coxon 1985: 182).

Estimates of relative movement between the Caribbean and North American plates provide a maximum value of 9.8 cm/yr (Mollet and Dinkelman 1972: 945). Caribbean (sea) records indicate an absence of large, continental polychaetes (the sole exception being the clath *Alvares* 1951; Morbido 1984; Valas and Orsiag 1978: 180), discussed later. The absence of other large Pleistocene mammals indicates that the main channels and passages of the Caribbean Sea remained open during the Quaternary and earlier.

Applying Milankovitch and Ewing's (1965) Atlantic sea level curves to the bathymetry of the Caribbean, it is possible to reconstruct the major outlines of the Carian-Caribbean region during the Holocene. Figure 5 presents a hypothetical reconstruction of the Caribbean by 17,000 B.P., at the end of the Wisconsin regression, with sea level 120-130 m (328-426 ft) lower than at present. By this time, Paleo-Indians may have been present in Central America. The record is rather poor, but there is evidence of limited



Figure 4. Tectonic setting of the Central Caribbean Region (after Molnar and Dinkelman 1973: 227, Fig. 3)



Figure 5. Hypothetical reconstruction of the Caribbean Archipelago for 17,000 B.P., with sea level 150-200 m below present level (after Molnar 1974a: 25, Fig. 2)

points from big game hunters in Costa Rica and Guatemala (Lyons 1978: 481). More importantly, it is assumed that Central America acted as a land bridge for Paleo-Indians migrating from North to South America. Unless we accept the radical theories of Amerind or transoceanic migrations (Catt 1978), it is clear that the human presence in Central America is older than the earliest South American Paleo-Indians, possibly dating back to 17,000–20,000 B.P. (Gale/South 1971).

According to Brown and Allaire (1978: 445), people started migrating into the West Indies possibly as early as 7,000 B.P. This entry date is supported by Pava (1971), Vokes and Griggs (1978: 149), Wiley (1976: 1), and others. Rapp (1978) suggests a possible entry between 15,000 and 6,000 B.P., while Katelson (1976a) suggests possible migrations as early as 13,000 B.P.

Regardless of the date of entry, it is clear that the first migrants used some type of watercraft to cross the passages from the mainland(s), and then from island to island. Prehistoric seafaring is discussed in detail later. The important point here is that the first West Indians did not walk onto any of the offshore islands. They were navigators in a water-endowed sea.

A lower sea level implies higher elevation of mountains and sea cliffs, says improved views of distant islands. Standing on a beach or bank, a man with 20/20 eyesight will not see a low-lying island 18 km (10 miles) away, for the curvature of the earth will hide it beneath the horizon of our hypothetical observer. If sea level were to drop sufficiently for that island to rise 34 m (110 ft) above sea level, then it would be visible from our hypothetical beach, 18 km away.

Paleobotanical and paleoclimatological studies indicate that much of northeastern South America was arid to semi-arid in glacial times, due to world-wide climatic compression (Lyons 1975: 440). In Central America and the Caribbean, the climate was cooler and drier, possibly resembling a Mediterranean climate (Nicholson 1976a: 28).

Radio-carbon, oxygen isotopes, and micropaleontological analyses of deep-sea cores indicate that, by 11,600 B.P., the Caribbean Sea was 1°C cooler than at present (Emmons et al. 1975: 1369). These cooler waters may have resulted from lower atmospheric temperatures in general, and from the discharge of glacial meltwater into the Gulf of Mexico via the Mississippi River (Lyons 1975: 440).

As the sea level fell, the hydrology of the Caribbean islands was modified towards equilibrium. Rivers cut into the "new" coastal plains of exposed tundra shelves, followed by downcutting of flood plains. Overall, river discharge was greatly reduced, while swamps and wetlands shrank due to the drier climate and to stream entrenchment (Jensen 1965: 13).

Cultural Transition

Let us suppose to imagine," said the Martian. "What does it matter who is First or Fourth? Let us both give the same flowers with equal frequency to a few thousand years. How do you know that those flowers are NOT the symbols of your own civilization into which I've collected them after a hundred and fifteen?"
 (Ray Bradbury (1947: 24))

We could have told our people the exact thing about the advancement of species, and again we would continue to be questioned further. But the Indian might ask, "Is it advancing, and toward what?" Or is it merely becoming complicated?"
 (John Barlow (1976: 271))

The Caribbean prehistoric complex covers cultural groups and ecological niches spanning thousands of years (Figures 4 and 5). The first

inhabitants of the archipelago were preceramic societies, arriving on the islands perhaps by 7,000 B.P. (Rouse and Allaire 1978: 440). But there were also preceramic (Archaic) cultures that survived into the so-called 'Ceramic Age' of Arawak-Carib peoples. Indeed, some preceramic groups survived into the early historic period, as reported by Spanish geographers Andrés Bernal and Alonso de Santa Cruz (Giffner 1958: 68).

Much Caribbean archaeological research has been concerned with chronology and taxonomy, at the expense of environmental reconstruction and sociocultural analysis. 'False' similarities between 'cultures' are common, generally resulting from the confusion of theoretical constructions with actual events in prehistory. Eventually, all cultural theories and taxonomies must be revised, refined, or replaced. Islands with known preceramic sites include Cuba (Falcon y Bay 1988), Hispaniola (Mason 1982, Rouse 1982, Velaz and Grijalva 1978), Puerto Rico (Grijalva et al. 1985, Pared 1978, Ota 1978, Vasquez (Figueroa 1978, Troncoso et al. 1984a), St. Thomas (Grove 1976, Telford 1978), Antigua (Hobbs 1978), St. Kitts (Rouse and Allaire 1978: 440) and Trinidad (Mason 1978). Few of these sites are much older than 3,000 B.P. If the hypotheses presented in this dissertation are correct, most early Caribbean preceramic sites are to be found beneath the sea.

Formal preceramic typologies have been produced by Kertész (1974), Flann et al. (1978), Rouse (1981), and Rouse and Allaire (1978). Most researchers would agree that there are three primary typological patterns, as summarized by Velaz and Grijalva (1978: 184). Pattern 1 is characterized by ground stone artifacts, including stone balls, mortars, and actual 'maces'. Pattern 2 is characterized by stone artifacts produced by percussion flaking, including choppers, scrapers, flint knives, and lamellae

Cuba	Yamalo	Yukl	Swampy Kapalik	Puerto Rico	Virgin Islands	Lower Antilles	1500s B.P.
1500s							1500
1600s		Yukl		Maricao (Yukl) Cajon (Yukl)			1600
1700s			El Financero	Yukl (Yukl) Cajon (Yukl)	Armed Bay	Yukl (Yukl)	1700
1800s		Cajon				Yukl (Yukl)	1800
1900s			Yukl			Yukl (Yukl)	1900
2000s			Yukl			Yukl (Yukl)	2000
2100s			Yukl			Yukl (Yukl)	2100

Figure 6: Prehistoric cultures of the Caribbean, including Trinidad (after Meggers et al. 1985, Orsini and Sauer 1988, Meggers and Sauer 1979, Pons 1971, Sauer and Allaire 1979).

Age	Cultural Unit	Cultural Typology	Series A-B
1500	Spanish, Indian, Dutch, African, American		1500
1600	Spanish to Early PR Yukl or Yukt PR	Spanish, Yukl	1600
1700	Spanish PR and W	Spanish, Yukl	1700
1800	Spanish, Spanish, Colombian	Spanish, Yukl	1800
1900	Spanish, Spanish PR and W	Spanish, Yukl	1900
2000	Spanish, Spanish PR and W	Spanish, Yukl	2000
2100	Spanish, Spanish PR and W	Spanish, Yukl	2100
2200	Spanish, Spanish PR and W	Spanish, Yukl	2200
2300	Spanish, Spanish PR and W	Spanish, Yukl	2300
2400	Spanish, Spanish PR and W	Spanish, Yukl	2400
2500	Spanish, Spanish PR and W	Spanish, Yukl	2500
2600	Spanish, Spanish PR and W	Spanish, Yukl	2600
2700	Spanish, Spanish PR and W	Spanish, Yukl	2700
2800	Spanish, Spanish PR and W	Spanish, Yukl	2800
2900	Spanish, Spanish PR and W	Spanish, Yukl	2900
3000	Spanish, Spanish PR and W	Spanish, Yukl	3000
3100	Spanish, Spanish PR and W	Spanish, Yukl	3100
3200	Spanish, Spanish PR and W	Spanish, Yukl	3200
3300	Spanish, Spanish PR and W	Spanish, Yukl	3300
3400	Spanish, Spanish PR and W	Spanish, Yukl	3400
3500	Spanish, Spanish PR and W	Spanish, Yukl	3500
3600	Spanish, Spanish PR and W	Spanish, Yukl	3600
3700	Spanish, Spanish PR and W	Spanish, Yukl	3700
3800	Spanish, Spanish PR and W	Spanish, Yukl	3800
3900	Spanish, Spanish PR and W	Spanish, Yukl	3900
4000	Spanish, Spanish PR and W	Spanish, Yukl	4000
4100	Spanish, Spanish PR and W	Spanish, Yukl	4100
4200	Spanish, Spanish PR and W	Spanish, Yukl	4200
4300	Spanish, Spanish PR and W	Spanish, Yukl	4300
4400	Spanish, Spanish PR and W	Spanish, Yukl	4400
4500	Spanish, Spanish PR and W	Spanish, Yukl	4500
4600	Spanish, Spanish PR and W	Spanish, Yukl	4600
4700	Spanish, Spanish PR and W	Spanish, Yukl	4700
4800	Spanish, Spanish PR and W	Spanish, Yukl	4800
4900	Spanish, Spanish PR and W	Spanish, Yukl	4900
5000	Spanish, Spanish PR and W	Spanish, Yukl	5000
5100	Spanish, Spanish PR and W	Spanish, Yukl	5100
5200	Spanish, Spanish PR and W	Spanish, Yukl	5200
5300	Spanish, Spanish PR and W	Spanish, Yukl	5300
5400	Spanish, Spanish PR and W	Spanish, Yukl	5400
5500	Spanish, Spanish PR and W	Spanish, Yukl	5500
5600	Spanish, Spanish PR and W	Spanish, Yukl	5600
5700	Spanish, Spanish PR and W	Spanish, Yukl	5700
5800	Spanish, Spanish PR and W	Spanish, Yukl	5800
5900	Spanish, Spanish PR and W	Spanish, Yukl	5900
6000	Spanish, Spanish PR and W	Spanish, Yukl	6000
6100	Spanish, Spanish PR and W	Spanish, Yukl	6100
6200	Spanish, Spanish PR and W	Spanish, Yukl	6200
6300	Spanish, Spanish PR and W	Spanish, Yukl	6300
6400	Spanish, Spanish PR and W	Spanish, Yukl	6400
6500	Spanish, Spanish PR and W	Spanish, Yukl	6500
6600	Spanish, Spanish PR and W	Spanish, Yukl	6600
6700	Spanish, Spanish PR and W	Spanish, Yukl	6700
6800	Spanish, Spanish PR and W	Spanish, Yukl	6800
6900	Spanish, Spanish PR and W	Spanish, Yukl	6900
7000	Spanish, Spanish PR and W	Spanish, Yukl	7000
7100	Spanish, Spanish PR and W	Spanish, Yukl	7100
7200	Spanish, Spanish PR and W	Spanish, Yukl	7200
7300	Spanish, Spanish PR and W	Spanish, Yukl	7300
7400	Spanish, Spanish PR and W	Spanish, Yukl	7400
7500	Spanish, Spanish PR and W	Spanish, Yukl	7500
7600	Spanish, Spanish PR and W	Spanish, Yukl	7600
7700	Spanish, Spanish PR and W	Spanish, Yukl	7700
7800	Spanish, Spanish PR and W	Spanish, Yukl	7800
7900	Spanish, Spanish PR and W	Spanish, Yukl	7900
8000	Spanish, Spanish PR and W	Spanish, Yukl	8000
8100	Spanish, Spanish PR and W	Spanish, Yukl	8100
8200	Spanish, Spanish PR and W	Spanish, Yukl	8200
8300	Spanish, Spanish PR and W	Spanish, Yukl	8300
8400	Spanish, Spanish PR and W	Spanish, Yukl	8400
8500	Spanish, Spanish PR and W	Spanish, Yukl	8500
8600	Spanish, Spanish PR and W	Spanish, Yukl	8600
8700	Spanish, Spanish PR and W	Spanish, Yukl	8700
8800	Spanish, Spanish PR and W	Spanish, Yukl	8800
8900	Spanish, Spanish PR and W	Spanish, Yukl	8900
9000	Spanish, Spanish PR and W	Spanish, Yukl	9000
9100	Spanish, Spanish PR and W	Spanish, Yukl	9100
9200	Spanish, Spanish PR and W	Spanish, Yukl	9200
9300	Spanish, Spanish PR and W	Spanish, Yukl	9300
9400	Spanish, Spanish PR and W	Spanish, Yukl	9400
9500	Spanish, Spanish PR and W	Spanish, Yukl	9500
9600	Spanish, Spanish PR and W	Spanish, Yukl	9600
9700	Spanish, Spanish PR and W	Spanish, Yukl	9700
9800	Spanish, Spanish PR and W	Spanish, Yukl	9800
9900	Spanish, Spanish PR and W	Spanish, Yukl	9900
10000	Spanish, Spanish PR and W	Spanish, Yukl	10000

Figure 7: Cultural Chronology for Puerto Rico and the Virgin Islands (after Sauer 1982, Sauer and Allaire 1979, Woodson 1978).

projectile points. Shell artifacts are absent, but some ground stone technology may be present. Finally, Pattern 3 is characterized by shell artifacts, including shell plectra, oyster vessels, and shell pagers.

Wetmore and Neupert

After the discovery of the masthead it was just a matter of time for the rest to go. One who was born by the water or had associated with Indians must be quite comfortable with it in any form.

John G. Wetmore (1916-1989)

*For a native to write on these subjects
 says an explorer's true friend. How fortunate
 George Wetmore, Lost Boyer (1988: 71)*

The enduring heritage of Caribbean aboriginals is dead. As Thomas (1953: 48) indicates for the aboriginal navigators of Oceania, the loss is "more than just a set of navigational techniques: it is the loss of a way of a life and a conception of the world."

Perhaps, seven or eight thousand years ago, a South American hunter and his family accidentally drifted on a raft down the Orinoco River and into the Caribbean Sea-but they do not concern us here. The permanent peopling of the Caribbean was a complex process involving countless social units and specialized cultural adaptations. Whatever the conditions of these first migratory waves, they certainly cannot be reduced to a customary Caribbean Family Robinson. To do so is to reduce a complex acculturational process to a mere historic accident.

The primary aboriginal method of water transport was the canoe, with possible supplementary use of rafts. The antiquity of canoe/making is well established. In Florida, prehistoric dugout canoes dating back to 3,000 B.P. have been found in wet sites (Bullen and Driskel 1997).

A raft without sails is a very poor craft for ocean navigation, and apparently rafts were introduced to the islands in later times. In 1606, a Free Indian, originally from Senegal, was rescued in Hispaniola by a British vessel. The runaway first narrated that he had been captured by Caribs, who eventually killed his companions, but spared him for teaching them how to rig sails to their canoes (McKean 1970: 8).

In the Bahamas, Columbus saw canoes 'all of one piece hollowed like a tray from the trunk of one tree . . . as large as to contain forty or forty five men, while others were as small as to hold one person' (McKean 1978: 8). In Jamaica, Columbus measured one finely decorated canoe at 28m (90 ft.) long and 3.6m (10 ft.) wide (Baker 1988: 82). This was surpassed by another canoe at Handford's Bay Islands, 'as great as a gallery, eight feet wide, all of a single trunk' (Baker 1988: 128). Ouyler (in Chittenden 1897: 143) describes a Lesser Antillean canoe 27m (90 ft.) long, which he estimated could carry 85 persons plus cargo. Columbus reports that the largest canoe traveled with great speed (Klein 1984: 32), but neither he nor any other early European explorer mentions the use of sails.

Mahogany (*Nicotiana glauca*), cedar (*Cedrela odorata* - Guttari), silk cotton (*Elaeis guineensis*) and other woods were used to produce oblong rafts (Makins 1976a: 30). First, a large tree was killed with fire and left standing to season, then it was felled and hollowed out with controlled fire, and the use of stone and shell picks, axes, and gouges. Shredded bark and polished, the log would then be soaked in water. Afterwards, the log would be inverted over a fire and heated. This would increase the wood's elasticity, so that the sides could be forced apart in order to give more beam to the dugout. 'Thousands of saws' must then be started, keeping the sides apart.

In the case of large canoes, the ends could be cut off prior to the final firing, allowing for maximum beam expansion. Afterwards, bow and stern pieces could be fitted with cordage passed through drilled holes, and the seams caulked with pitch from various plants (Nicholson 1976: 308). The hull could be caulked with a number of mixtures. The *Questione di Ecuador* for instance, has a mix of shell powder and beer (Gervasio 1987). An anonymous French chronicler indicates that the Caribs painted the stern of their canoes in red, using a 'red earth' which was supposed to be the feces of a large snake called *Ochato* (in Christian 1861: 188). Reinhardt (in Christian 1981: 378) reports that the Caribs at times painted mythical creatures on the foreboard before the water line of their canoes.

Prehistoric Caribbean designs had two types of dogouts: the canoe and the pirogue. Pirogues generally refer to the larger canoes used for inter island voyaging or warfare. But one alone does not define a pirogue. The pirogue has a relatively wide beam, achieved through the process of controlled heating described above. Two other important traits of the pirogue are the use of wooden planks to increase the foreboard, and the presence of a keel for increased stability and control (Smith 1983: 345).

In Coahuila, native boat builders achieve additional stability by the use of an outrigger (Ballester 1979: 390). For increasing both stability and cargo space, two pirogues may be attached by two or more light beams, forming a double pirogue (Ballester 1974: 432). Apparently, these last two techniques had limited or no use in the prehistoric Caribbean, as they are more typical of sailing canoes.

Lewis (1938: 427) suggests that the prehistoric Tainos of the Greater Antilles had canoes, while the Caribs of the Lesser Antilles had pirogues. However, as McEwen (1979: 83) points out, Lewis' reference to

Carib navigation was de la Bode, an 18th century French author. By that time, Carib navigation had been greatly influenced by Europeans, including the use of sails, as reported by Labat (Caribbees 1360-580).

At this point, there is no reason to assume that the prehistoric peoples were any less adept than the Tainos or Caribs of protohistoric times. If anything, the prehistoric cultures may have been even more maritime-oriented than later agricultural peoples. As domestic animals, sails were probably introduced into the Caribbean at historic times, but canoe-building is an ancient art.

The building of a prehistoric canoe must have been a special thing, full of ritual and symbolic meaning, just as Pele in Caribbees (1891: 312) observed for the Caribs in the 17th century. Nicholson (1896: 300) is right on the mark by suggesting that the canoees sailed for much "diversion and merriment." A deep, near mystical feeling for the construction of watercraft is shared by all "primitive" boat builders, a feeling that is universally understood by mariners anywhere.

McKeanek (1876: 7) questions the feasibility of carving a 28m (90 ft.) long canoe out of a single log. But here he is wrongly imposing modern, limited conditions on the prehistoric environment. Seven thousand years ago, tall mangrove and cedar trees were commonplace both in the islands and on the mainland.

In the early 1880s, a group of kayakers attempted to paddle from South America to Florida, following the Caribbean island arc. At Mona Island, west of Puerto Rico, they reported that their only serious problem had been the occasional encounter with large pelagic sharks (Arthur Combs, personal communication 1946). They had no problems with winds or currents, and fortunately no storms were encountered at sea.

In 1987, four 15m (49 ft.) long dugout canoes departed from the Andes to embark a commemorative journey through the Amazon, and up the Caribbean islands to San Salvador in the Bahamas (Hanes 1987). In the winter of that same year, one of the canoes sank off Vieques Island and was promptly salvaged by local divers. Weeks later, while conducting the marine archaeological survey for the Vieques-Guadalupe water pipeline, the author had the opportunity to inspect the canoe (Wynn 1988).

Contrary to the tenets of experimental archaeology, iron nails and screws were used for thwarts attachments. The structurally loose seats and the narrow beam indicated that beam expansion had not been employed, as confirmed by a description of the building process published in a Spanish newspaper (Correos de 1887: 28). Constructed by Quechua Indians in Ecuador, this was a river canoe.

In Vieques, the author's inspection revealed no hull damage resulting from collisions or torpedos (Florida *torpedos*) worms. The canoe simply floundered. Was this unfortunate accident solely the result of rough seas, possibly combined with poor seamanship? Or was the canoe not entirely suitable for offshore navigation?

It is possible that the pirogue is a genuine aboriginal invention, an adaptation of river canoes to oceanic conditions. Beam expansion would make a more stable canoe. Floating low in the water, the canoe going ocean would require planing to reduce water intake. A semi hull or perhaps simply a hard floor could be carved out of the original log, for added stability and control. Of course a full hull only makes sense in a sailing craft, as in the sharp hulls of Austronesian outriggers (Ballard 1973: 108, Figures 11.1 and 11.3), but a hard base would seriously improve the stability of a human-powered canoe.

It is interesting to notice that the Puerto Rican canoe race, traditionally carried in a pointed paper cup, is called *peragua*, which is Spanish for *peragua*. Of course *peragua* here is not a Caribbean prehensile trout, but the word *peragua* is. At face value, it seems quite a jump from linguistics to boat construction, yet, the relationship between the word and the paper cup is not entirely fortuitous, for in the age of sail a *peragua* was a beaked (pointed hull) canoe (Gale 1885, Tolson 1909, Voss 1990). Did the word have the same meaning in prehistory? If *peragua* simply means canoe, why would the aborigines also have the word *hawa*? Two words for the same object?

According to Coll y Toste (1878: 246), the word *peragua* is a South American term for small boat, deriving from the root *para* which in Guaraní means fish. In a more sophisticated study, Adams (1877: 102) concluded that *peragua* is a Tano word. This is supported by Taylor (1877: 20), who points out that the Tano word for water is *hagua*, or *hikaua*, while the word for God is *Tikoua*. *Agua* (Maroon), roughly translated as Lord of the Maroon and the Sea, Who has no Father (Patt 1974: 21, 27).

In the author's opinion, the term *hawa* was used by the Tano both as a generic term for all watercraft, and for designating the round-hulled canoe used for river and inshore navigation. On the other hand, the term *peragua* was used to designate a large, planked, possibly semi-beaked canoe for ocean navigation.

In the final analysis, the question of whether *hawa* or *peragua* were used by the prehistoric navigators can only be answered unambiguously, if we ever get lucky enough to find and properly interpret an ancient aboriginal canoe. To date, no aboriginal canoes have

been found in the Caribbean, not even perichthene was. Sharing the fate of numerous Maya volumes, Taina records were often burned by the Spanish authorities in order to suppress aboriginal travel.

The maritime orientation of Caribbean aboriginal culture is confirmed by virtually all the European chroniclers from Columbus onwards. Numerous historic narratives speak of the great swimming, diving, fishing, and maritime-hunting ability of the aboriginals. For instance, Lohel de Gárdano (1881-1890) describes how a harbor sank off Martinique in 1489 with everyone drowning save for an Indian who stayed afloat for two and a half days without any support. In 1576, a large *poisson-fur*, from the description clearly a great hammerhead shark (*Sphyrna*), fatally wounded a boy at Nassauville, St. Kitts. Lohel narrates how a Dutch Britonnet dove and killed the shark with bayonets, tied a rope to its tail, and towed the "monster" to the beach, where the dead boy's leg was retrieved from the shark's stomach.

Cannae were the single most important objects in the prehistoric occupation of the Caribbean, and it is deeply symbolic that they may also have been an important encouragement for the historic conquest as well. In a fascinating and seldom discussed passage, de Las Casas (1890: 47-48) indicates that aboriginal natives as captives drifted to the Azores, Madeira and Canary Islands, adding credence to the mysterious legends of lands beyond the great Ocean Sea.

Among the sailors who reported on these mysterious "isls" were Pedro Quares, Columbus brother-in-law, who was one at Ponta Santa (Madeira Islands), and Martin Wascia, who plotted one up 400 leagues (1600 miles) west of Portugal's Cape St. Vincent. De las Casas describes some of these lands as modern islands (worked or carved logs), but he

also refers to them as *cannoes* and *climades*, the latter being the old Hispanic Indian word used prior to the inception of the Caribbean term *canoe* (Gusley 1981: 35). There is no doubt that these were New World canoes.

Cultural Ecology

We do not think a fully-man-constructed shelter, but good that, we had a spot, we would be there daily to find, about it, and thought. And a nature of my contemplation there would be, I suppose of looking from across that very nature were attracted.

John Cheever (1976: 198)

But it is good that we do not have to fly to all the suns in the crown of the moon. It is enough to live on the sea and think for the future.

Robert Heinlein (1968: 47)

The Pleistocene-Holocene transition was a time of great ecological and cultural transformations. Due to either climatic changes or human over kill, or both, numerous species of large land mammals became extinct. In the Americas, this megafauna included mastodons, mammoths, megatherium, mylodon, peledonia, camels, hippobius, toxodon, river bears, saber-toothed tigers, dire wolf, glyptodons, etc. Many of these animals were already extinct or nearly so by the end of the Pleistocene, some persisted until the mid-Holocene (Dickens 1988: 136, Lynch 1975: 478).

The archaeological record indicates that by 8000 B.P., New World prehistoric cultures intensified their reliance on gathering and exploitation of coastal resources (Cohen 1987: 281). By 7000-6000 B.P. (if not earlier), permanent settlements appeared on estuarine environments (McNeish 1967: 317). Overall, the early Holocene was a time of relatively rapid sea level rise, in the order of 30 cm (1 ft) per year (Kennedy, 1982: 278). Bappel (1990a: 38) has suggested that coastal sites were formed during *stillstands*, periods of a temporary halt in sea level fluctuations. In times of

rapidly rising sea level, coastal villages may have formed on a semi-permanent basis, either moving progressively landward or to new sites nearby. Certainly this would have had a tremendous impact on site formation, in terms of midden size, stratigraphy, and horizontal displacement.

The technology of water-crafting allowed a maritime-hunting lifestyle as complex as ritual and technique as any land-based, Pleistocene hunting tradition. The marine prey of Central-Caribbean hunters included manatees, seals, sea turtles, sharks, rays, and stranded whales. Coastal fauna included birds, iguanas, small mammals, tortoises, and shellfish (Tighe 1978, Olsen 1992). A variety of wild plants and fresh and salt water fish would complete the diet (Wing 1990).

The maritime hunter would have been characterized by two primary tendencies. The first tendency would be to move when the animals they hunted became scarce, thus the entry of prehistoric peoples into the Caribbean. The second tendency would be a progressive move from marine big game to small coastal game, and a further specialization of shellfish collecting (Turner 1980: 104). This second tendency may have led the coastal hunters to greater permanency in their settlements.

The primary prey may have been the manatee, a large mammal found in both fresh and salt water. Attaining a length of 4.2 m (14 ft), manatees were hunted in prehistoric and precolonial times throughout the Central-Caribbean, Florida, and the Gulf of Mexico. Manatees belong to the order of *Sirenia*, comprising five modern species of aquatic, herbivorous mammals: 1) the dugong (*Dugong dugon*), ranging from the east coast of Africa to Southeast Asia, Australia, and the Solomon Islands, the African manatee (*Trichechus senegalensis*), found in the coasts and rivers of West

Adria, from Senegal to Angola, 2) the Amazon manatee (*Trichechus inunguis*), found in the Amazon and other South American rivers, 3) the West Indian manatee (*Trichechus manatus*), found in Central America, the West Indies, and Florida, which some researchers divide into a Caribbean and a Florida subspecies, and 4) Steller's manatee (*Hydrodamalis pastus*), possibly hunted to extinction, although there are reports of survival (Barrett et al. 1979: 45).

The author has discussed manatee biology, ecology, and preference hunting techniques elsewhere (Vega 1981, 1985), and will only summarize here the hunting data. For an extended bibliography on Steller, see Warfield and Farnington (1975).

Stellerians are endangered everywhere, but some traditional hunting continues to hunt them as their primary source of meat. In Australia, a harpoon and rope are still used to hunt the dugong (Hemmer 1979: 49). The same method has been used into modern times by the Makoto Indians of Nicaragua (Barrett 1955: 217). The Maja of Ecuador also used a harpoon and line for manatee hunting (Baughman 1948: 297). The Guayana Indians hunted the Amazon manatee with shell-tipped harpoons (Stearns 1979: 178). Manatee hunters of the Orinoco River used a double-barbed harpoon tied to their waist with a strong rope of manatee hide (Baughman 1948: 297).

In the Comayma River, between Guyana and Surinam, Amerind Indians used three-pronged arrows to hunt manatees. On Nicaragua's Indio River, manatees are traditionally hunted at night (Barrett 1955: 219). Some Amerindians throughout the area use nets to capture dugongs (Colley 1977: 284). In Guyana, an entire Indian village may engage in constructing underwater stockades to trap manatees and turtles in strings over

channels, drowning captured manatees by placing wooden pegs in their nostrils (Wing 1961: 3). Pedro Martir (in Arroyo 1972, Underhill 1982) reports of a Greater Antillean Taino chief who netted a young manatee and named it as a pet.

José González, a Jewish missionary who lived among the Indians of the Orinoco River in the late 18th Century, describes how hunters would transport a dead manatee. Leaping into the water, they would tilt their canoe in order to flood it, then they would push it under the dead manatee, hold out the water and paddle home (Broughman 1948: 277). All sources report the use of canoes. Excluding river environments where nets could be worked or man screens, the canoe was an indispensable tool for manatee hunting.

Manatee bones have been found in archaeological context throughout the Caribbean, including the sites of Colmanes III, Cuba (Gervasio 1968: 133), Sabana de Juanillo and El Perreño/Serralito in the Dominican Republic (Valer and Orsaga 1976: 151, 160), Anacoí, Dominican Republic (Rosen 1968: 196), Cueva María de la Cruz, Puerto Rico (Salgado et al. 1982: 117), Culebras, Las Cañetas, Pinarque and Playa Blanca, Puerto Rico (Rosen 1968: 207, 264, 261), Isla Verde, Puerto Rico (Vega 1981: 52), many sites in Guadalupe (Wing 1968: 100), Chet, St. Lucia (Ray 1969: 412), etc. Charlevoix (in Perles 1907: 56) mentions a Haitian canoeist killed by the nose of a manatee.

Observations of manatees in lagoons have been made by Columbus (Coffe 1964: 111), Ovando (in Rosen 1968: 54), Martir and Ovando (in Arroyo 1972: 52), Carvajal, de Leanda, Alonso, González, Raleigh, Harcourt, Tempier, Espenewick, Clarys, Humbolt, William Bates (in Broughman 1948), and Labat (in Ray 1969: 413).

When the first pelagosian sealions reached the West Indies, large manatee herds were commonplace. In 1883, a diving herd encountered in Morston Bay, Australia, reportedly extended over an area of 8 km (5 miles) long and 274 m (300 yards) wide (Hazardus et al. 1978). Unquestionably, similar manatee herds once lived in Caribbean waters. Before observed years ago, the manatee was to the Caribbean what the elephant was to Africa and the bison to North America.

Another important prey of Caribbean marine hunters was the sea turtle. Five species were available: 1) the loggerhead turtle (*Caretta caretta* L.), 2) the Atlantic Ridley or Bandier turtle (*Lepidochelys* sp.), the Hawksbill turtle (*Eretmochelys imbricata* L.), 4) the green turtle (*Chelonia mydas* L.) and 5) the leatherback or trunk turtle (*Dermochelys coriacea* L.). Turtle bones have been found in numerous archaeological sites, including Pontalva and Caba San Antonio, Cuba (Raussey 1948: 146), Gamañá Bay Caves, Dominican Republic (Raussey 1948: 125), Las Cochinos, Puerto Rico (Raussey 1958: 347), Cape Cofresi, Puerto Rico (Maggiolo et al. 1978: 20), Rincón de Honda, Puerto Rico (Tronolone et al. 1984: 201), Isla Verde, Puerto Rico (Vega 1961: 56), Magasa Bay, St. John (Majumdar 1962: 46), etc.

In his 1842 expedition to the Amazon, Francisco de Orellana observed that the Orongo Indians "raised thousands of turtles in ponds beside each house" (Hemming 1978: 14). In Surinam, the modern Matsigena men used to capture sea turtles, but the traditional method was to hunt them with wooden shafts (Matsigena 1977: 87). Turtles could also be located with bow and arrow. The Tsimé of the Greater Antilles used an ingenious system of tying a line to a ramrod or waterfowl (*Ardeotis* sp.), letting it attach to the turtle's plastron or carapace, and then pulling both back to the canoe (Raussey 1948: 81, 150). Finally, turtles could be captured manually

incubated on the beach, as they came ashore to lay their eggs (Wing 1948: 190).

Stranded whales and porpoises were a potential source of meat to prehistoric hunters. The use of stranded whales in the European Middle Ages has been recorded for Tardusians, Larmans, Chiansin, and Bristol's cultures (Cohen 1977: 119). Humpback whales (*Megaptera americana*) appear regularly in the Caribbean from January through March (Edwards et al. 1978: 8). These migrations are necessary for survival, for the young calves lack the blubber thickness necessary to withstand the icy Arctic and Antarctic waters (Edwards 1971: 88). Undoubtedly, the colder climate of the late Pleistocene-early Holocene resulted in an increased Callosia presence in the Caribbean.

Other whales that are regularly seen in Caribbean waters include humpbacks or fin whales (*Megaptera* sp.), Cuvier's beaked whales (*Ziphius cavirostris*), pilot whales (*Globicephala macrorhynchus*), and occasionally sperm whales or melonates (*Physeter catodon*). Bottlenose dolphins (*Tursiops truncatus*) and common dolphins (*Delphinus delphis*) are frequently seen, with occasional sightings of Pacific dolphins (*Stenopus gamma*), spotted dolphins (*Stenalia* sp.) and killer whales (*Orcinus orca*). In 1975, a pack of twenty-five killer whales or orcas attacked a large whale near Grenada Shoals, SE of Cuba's Island (Edwards 1979: 836).

Whale strandings are relatively common in the Caribbean. In 1981, a beaked whale was stranded off La Parguera, on the south coast of Puerto Rico, followed by five strandings in 1985. Another beaked whale was stranded in 1984 at Petilla Beach, east of Ponce (Edwards 1978: 837). In 1988, a sperm whale was stranded off Punta Cheloa, west coast of Puerto

Race, that same year, another sperm whale was stranded near Villa del Mar, Dominican Republic (Bracegirdle et al. 1973: 4).

The use of whales by Caribbean prehistoric hunters is proven by a mortuary assemblage found at El Perreño/Barahona, Dominican Republic. The deposit also contained associates, sea turtles (*Chelonia mydas* L. various species of fish, and shellfish, ray (*Apeltesmelus* sp. L. turtle (*Phrynobolina* sp. L. and possibly crocodile (Pérez and Ortega 1976: 154).

In the 18th Century, New England whaling vessels regularly hunted sperm whales, pilot whales, and humpbacks in Caribbean waters. Late in the century, at least ten native whaling stations arose in the islands. Using small, open vedettes 7 ftm (22-26 ft.) long, the whalers would chase their prey with "wrought iron harpoons and "death lances," an explosive lance splinter fired from a shoulder gun (Graham 1971).

It is extremely unlikely that Caribbean prehistoric hunters actually set out in pursuit of whales, but then again every age has its deadends. A far more realistic scenario is the hunters taking advantage of stranded whales, killing the animals with their pointed spears such as those of the Caribbean Caribbeo complex of the Dominican Republic (Pérez and Ortega 1976: 163, Figures 90-93).

The prehistoric aborigines probably hunted the West Indian monk seal (*Monachus tropicalis* Gray), which is historic times ranged throughout the Caribbean Sea and the Gulf of Mexico (Colley 1977: 190). Although generally regarded as extinct, apparently a few individuals have been sighted in recent years (Slate 1988: 78).

From the above discussion, it is clear that the Caribbean and other tropical areas of the world had the resources for a maritime hunting tradition. Just as there were Pleistocene big game hunters on the

mainland, there may well have been early and *Heliconia* big-purse buntings on Caribbean waters. These maritime buntings may have been the first settlers of the Antilles.

Slacks were the largest Pleistocene land birds to inhabit the Greater Antilles. According to Morrison (in Yule and Ortega 1978: 197), eight species of slacks lived in Cuba, six in Hispaniola, and two in Puerto Rico, including the genera *Megapodiceps*, *Megascopus*, *Parus*, and *Arremonops*. Slacks were common in the Americas mainland as early as the Pliocene (2 to 15 million years B.P.), which suggests a Caribbean migration well before buntings. For Rugg (1973: 158) and Korbis (1964: 29), the presence of slacks extant skeletons suggests a Caribbean land bridge with Florida or Honduras. But if a land bridge existed, how do we account for the lack of other large Pleistocene megafauna in the Greater Antilles? For instance, in Trinidad, which was connected to South America until 10,000 B.P., mastodon fossils have been found at Pitch Lake, Forest Reserve, and Los Bajos (Morris 1976). Similar fossils have never been found past the Tehuantepec Isthmus.

Kempson (in Norris 1979: 3) has suggested that the indigenous fauna of the Greater Antilles is the product of continental rafting from the Magdalena and Orinoco Rivers, in Colombia and Venezuela respectively. If we take into consideration the prevailing NEW currents of the Caribbean, this hypothesis explains the presence of slacks in the Greater Antilles, and their absence in the offshore Lesser Antilles. Although some species of slacks were fairly large animals, it is quite possible that they would be rafted during odd tide downriver. On the other hand, it is virtually impossible that mastodons would be transported in similar fashion.

The migration routes of island prehistoric cultures is a mystery, and it will continue to be so until sufficient early sites are available to generalize. If the hypotheses presented in this dissertation are correct, then most of the early sites that we need to find are beneath the sea.

No completely submerged prehistoric sites have been excavated in the Caribbean, and apparently none have been located. So far, Caribbean prehistoric sites have been found on uplifted or prograding coasts, or on small islands and keys separated from the parent island, where the island workshops with only an indirect relationship to the coast. The real prehistoric sites, some of which are close to or within the intertidal zone (Yales et al. 1978). It is no coincidence that many of the early prehistoric sites have been found in Hispaniola (Ortega and Guzman 1981, Yates and Ortega 1978), an island that has experienced localized uplift well into the Pleistocene (Alexander 1968: 181).

South, Central and North American origins have been proposed for the prehistoric peopling of the Caribbean (Ortiz 1968; Cronin and Brown 1968; Flenk 1971; Rapp 1973; Brown 1964, etc.). Bullen (1975: 14) has argued against a Florida origin on the basis of the Gulf Stream. The Stream segment that separates Florida from Cuba and the Bahamas is the Florida Current, with a velocity of five knots (nph), and a transport rate of about 20 million cubic meters of water per second. Further north, the Stream runs somewhat slower, but the transport rate increases to over 100 million cubic meters of water per second, roughly 100 times the flow of all the rivers on the world put together (McLain 1969: 47).

Certainly the Gulf Stream is an enormous "river in the sea," but it would be a mistake to see it as an insurmountable barrier to southward prehistoric migrations. First, a lower sea level would reduce both the

volume and speed of the currents, second, the greater size of the Bahamas Islands would reduce the crossing by a half in some places. Contrary to the usual argument, the easier crossing would not be from the Florida Keys to Cuba, but rather from East Florida to the Little Bahamas Bank, beneath the 27th parallel. Once in the Bahamas, the pelagians would have easy access to Central Cuba and Hispaniola via Turke and Caima, Ketchikan Bank-Silver Bank. Without any intention of participating in the "postwar's waste game," it is worth mentioning that a number of submerged, early sites have been discovered in the Vero Beach area, East Florida (Cockrell 1990). At present, the author is not inclined towards any route in particular. Few early sites are known, and Caribbean paleoenvironmental reconstruction is in its infancy, as is the cultural ecology of tropical, maritime hunter-gatherers (Tanner 1993: 746).

CHAPTER 4 COASTAL GEOARCHAEOLOGY, PUERTO RICO

Physiography and Climate

After one has lived in those latitudes long enough the changes of the seasons become as important there as anywhere else and Italian sailors, who loved the coast, did not want to miss any spring (the summer), nor any fall or winter.

(Columbian History [1492]: 9)

There is a passage in the pathless woods

There is a harbor on the rocky shore

There is a noisy wilderness of vines,

By the steep hill, and through its cave

(George Gordon, *Lord Byron* [1808]: 74)

Puerto Rico is approximately 341 km (210 miles) long and 54 km (33 miles) wide. The island is divided into a northern and southern half by a central mountain chain of volcanic and plutonic igneous rocks (Olschoff 1954: 80). About 40% of Puerto Rico is mountain terrain, with 30% hill, and 30% level (Spot 1974: 24). The highest elevation is that of Cerro Punta, towards the center of the island, with an altitude of 1,038 m (3,398 ft). Prominent peaks also rise on the northeast section of the island, at the Luquillo Range, reaching a maximum elevation of 1,174 m (3,833 ft).

The regular shelf is narrow to the north, and wide to the east and southeast (Figure 8). The composition of the coastline ranges from unconsolidated sediments to limestone and granite formations. Of 740 km of coast, 80% are beaches, which are generally short and divided into separate systems with little interaction (Olschoff and Truesdell 1955: 147).

Tides average about 30 cm (1 foot). Shoreline exhibits great variation, with five generalized types in six separate stretches (Kaye 1988: 51).

Sea surface temperatures range from 32 °F in September to 77 °F in February. Largely due to the moderating effect of the sea, Puerto Rico's median temperature of 79° at sea level varies within 3 °F throughout the year (McDowell 1988: 15). Most rainfall and river discharge is to the north, with a semi-arid north coast. Rain is most abundant at El Yunque Rain Forest on the northeast, with over 8 in (204 inches) of annual rainfall. The southeast is the driest part of the island, with an average annual rainfall of 1 in (142 inches).

Puerto Rico's position on the northeast corner of the Caribbean Basin exposes it to the mainstem of the Great Northern Equatorial Current (GNEC). This powerful ocean current originates off the West African coast and crosses the Atlantic in clockwise fashion, veering north as it reaches the Caribbean. The influence of the GNEC provides an additional element of humidity to the island's tropical climate. The on-land breezes generally blow offshore at night and at dawn (when the land is cooler than the surrounding ocean), and onshore during the daytime (when the sea is cooler). The island is also affected by hurricane winds, which are frequent in the Caribbean during the warmest months, from June to November. In general, hurricanes follow the SWW path of the trade winds. The coastal plains are narrow and flanked by hills. The central mountain ridge slopes gradually to the north, and steeply to the south.

There are approximately 1200 streams, of which approximately 110 form true river systems. The largest rivers flow to the north, where low drainage results in swamps and marsh lagoons. On the interior of the island are numerous wet and dry river systems. Soils show a high degree of

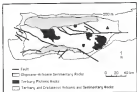


Figure 8. Tectonic setting of Puerto Rico (after Morelock 1978 & Fig. 11, Morelock and Trumbull 1980: 348-50, Fig. 20-1)



Figure 9. Location of profile lines used for the study

variation, with only about 6% or 125,000 acres classified as first rate. At present, about 70% of the island's soils are considered of inferior quality, located mostly in the mountainous interior (Paul 1974: 214).

Puerto Rico is an excellent setting for the study of sea levels. As Kays (1956: 51) has pointed out, the island is "exceptionally well suited" for the study of saltwater and beachrock. These lithified sand formations are both resistant and highly visible, and may be used to trace the exact location of recent shorelines (Kays 1956: 47). Moreover, saltwater and beachrock provide accurate "signposts" of sea level in the form of rips and tidal terraces, and may be radiocarbon dated (McLean 1967, Williams and Emery 1968).

The island has a wide variety of coastal geomorphological settings, yet is small enough to be sampled as a single project. The north coast is characterized by high-energy beaches and limestone cliffs, in contrast to the south a low-energy, predominantly mangrove coastline. Both the east and west coasts exhibit fringing coastal (Meredith and Trumbull 1955, Polunin 1978). A rich land topography provides additional evidence of marine transgression in the form of submerged speleothems.

Puerto Rico has a long prehistory characterized by coastal culture and one of the longest historic periods in the New World, with rich documentation and well-preserved constructions of coastal installations (Gibson 1980, Lopez Camps 1975, Via Ybar 1974). Early cartographic and architectural material abounds, including nautical charts and fortification profiles with marks sea level.

The small tidal range of 30 cm (1 foot) reduces the inter-tidal zone to a very thin zone pinned to a narrow strip, somewhat simplifying the archaeological measurement of coastal change. Finally, the island's

geomorphology is not further complicated by periglacial rebound, and adequate to excellent driving conditions are found nearly year-round.

Geomorphological Markers

It is advisable to look from the hills past to the sea, and then back to the litiged
again.

John Galsbrook (1876, 1914)

Today I sit under from the clouds where men once gathered from the sun.
Michael O. Flannery (1875, 189)

The Caribbean is characterized by a number of geological formations closely related to sea level. A complete listing of these features is provided in Chapter 1 (Research Summary). In the course of field work, the author has had the opportunity to make extensive observations of some of these geomorphological markers of coastal change, particularly eskante and headbreak.

The coastal geomorphology of Puerto Rico is formally discussed by Gades and Glass (1957), Kaps (1959a, 1959b), Mitchell (1954a, 1954b), Morlock (1973), Morlock and Thornhill (1983), and others. We shall now take a brief look at some geomorphological considerations pertinent to the study.

Eskante

Eskante is wind-deposited sand cemented by calcium carbonate (CaCO_3). In Puerto Rico, these cemented or lithified dunes are prevalent along the north shore, in sections reaching a height of 30 m (100 ft.) or

more. San Juan is impressive between of St. Maria and San Cristobal rose along volcanic, of which four generations of cemented dunes are recognizable (St. Rd), each separated by a layer of ancient sands or paleosols (Rape 1984a: 18). The white sand ridges off the north coast, from Camuy to Lajas, are the remnants of submerged remained dunes of late Pleistocene origin (Rape 1984a: 83, Polley 1971: 99). Many of Puerto Rico's coral gardens exposed by the water are not true coral reefs, but rather patches of submerged volcanic covered by a thin veneer of coral.

Cemented dunes are weathered by a complex process involving frontal attack by waves, undermining of sand beneath the cemented dune by sand, the waving and boring habits of periwinkles (*Littorina*, *Littorina*, *Littorina* and *Tectarius* *maritima* Linné), barnacles (*Balanus* *balanus* L.), clionids (*Cliona* *intertextilis* L.), and sea urchins (*Diadema* *diversum* L.), biochemical activity as tide pools, and the milling action of surf-propelled pebbles.

Looking at the distribution of cemented dunes in natural charts and aerial photographs, it is clear that much of Puerto Rico's north coast has evolved from a relatively straight volcanic shoreline to an eroded coast. This is the product of erosive wave patterns deflected by breaks in the volcanic ridge, an important process for understanding the importance of at least one archaeological site, that of Isla Verde (Rape 1981: 14), discussed in more detail below.

Driving half a mile or so off the north coast, the water has seen cemented dunes starting at 1 m (33 ft.) beneath the surface of the ocean and protruding to a depth of over 40 m (131 ft.). These submerged volcanic formations mark a Quaternary paleoshoreline far less explored than the Atlantic-ocean coastline, remains unknown.

Beachrock

Beachrock is cemented beach sand or shingle. As is evident, the cementing agent is calcium carbonate, although rare instances of cementation by iron oxides have been reported by Kays (1959b: 60). Beachrock forms a hard pavement along the shore (littoralized), when found underwater, it clearly indicates that marine transgression has taken place. The pavement generally dips at a slight angle towards the sea – and may reach a width of 60 m (200 ft.) or more. Trails or grooves perpendicular to the shore are often formed on the pavement, produced by mechanical stresses (McLean 1967). Like coral, beachrock is a formation of tropical and sub-tropical seas. The distribution of beachrock formations in Puerto Rico is discussed by Gilson and Glass (1967) and Kays (1959b).

Off Puerto Rico, numerous submerged beachrock pavements may be observed by scuba and trawl diving. Off Ship-Is., at a popular surfing spot just west of El Estero de Sports Complex, San Juan, the author has traced an underwater beachrock pavement that runs for over 200 m (1,300 ft.). Kays (1959b: 118, Figure 59) also reports on it. While conducting the Virgin Island Islands Archaeological Survey, the author found a beachrock pavement at a depth of 15 to 60 ft., four miles from the nearest land (Kays 1958: 60).

Sea Caves and Rock Caves

Sea caves are formed by wave action on high energy coasts, and finding them underwater clearly indicates that marine transgression has taken place. The author has explored sea caves off Costa Azul (Laquele), Bajura de las Cervezas (Sancti Spiritus), Isla Verde, Culebra

Island, and Moss Island. Other divers have reported unknown caves at Toumaline Reef (off Mityaga), La Furguera, Niagova Island, etc.

Karst caves are formed by the weathering and erosion processes of acid-charged water on soluble-rock terraces, including limestone, dolomite, marble, and gypsum. Puerto Rico has a impressive karst topography, with an estimated 2,000 caves and at least one 12 km long system at Caguas (Adams 1993, Goffe 1993, Morán 1993). At Guayama Beach, the author discovered a small karst cave that is now beneath the sea. In the mouth of two veins, which should only be explored by experienced cave divers in the calmest sea, fish swim amid short stalagmites and columns.

Archaeological Sites

"You always worked in water, Martin," said Michael. "Where are they, Dad? You drowned."

"Drowning was," said Dad, with the author Michael on his shoulder and pointed straight down.

The Martins were there. Timothy hopes to return.

The Martins were there—in the waves—reflected in the water. Timothy and Michael and Robert and Alan and Jack. The Martins almost lost it then for a long long time less than the dying water.

Ray Bradbury (c.1967-1975)

On white shell and sand and foam

On the margin of the stones,

Where a low grey fungus grows,

Roundly of the sea and the land

P. B. Shreeve (1976-1987)

The coast of Puerto Rico is rich in submerged sites. All four coasts and all large offshore islands possess evidence of prehistoric occupation. Numerous archaeological studies have been undertaken since the pioneer work of the late 18th and early 20th centuries, including the early research by Fiske (1897) and de Monts (1919), and the extensive archaeological surveys by Bailey (1940) and Rowe (1939a, 1939b).

Following a literature search and interviews with various archaeologists and coastal dwellers, beach and marine surveys were conducted by the author along Puerto Rico's four coasts. Over 100 marked and unmarked sites were evaluated, ranging from solo dives to groups of up to six divers, generally following transect lines perpendicular to the shore. The length and spacing of transects varied according to variability and bottom contour. Three simple operations included the use of underwater slates, measuring tapes, PVC stake rods, compass, and underwater cameras and video. In the case of Isla Verde Site, a more complex field strategy was involved, including underwater navigation with a water diver and PVC grids, as well as aerial photography analysis (Yago 1981).

Coastal and submerged sites were mapped onto the U.S. Geological Survey Quadrangles using the Puerto Rico 2000 meter grid coordinate system. Due to the nature of the research, the standard presentation of contour intervals in meters or water depths in feet was a subtle visual hindrance to the blending of submerged and subaerial geomorphological interpretation.

In some cases, access to beaches proved to be a logistical problem. As an inheritance from Spanish law, the beaches of Puerto Rico are public lands (Costal Trusts 1972: 28). However, due to the nature of Puerto Rico's beaches—which are generally small and segmented by headlands, lagoons and rivers, access often requires passing through private lands, or the use of a boat. Today, Puerto Rico's coastline is segmented by harbors and *Mis Dique* signs. On occasion, the harbors cut across the beach and into the sea, in violation of the law. Much of the coastal zone outside the metropolitan area belongs to absentee owners, further complicating the process of acquiring permits for archaeological surveying.

A total of twenty-one prehistoric coastal sites in Puerto Rico were chosen for the study (Figure 6). In terms of location, the sites range from over 1 km inland to submerged (Table 2). Many of the sites were surveyed or visited by the author, generally including restricted surface collecting. Most of these sites pertain to the Ortoiroid series or culture, which has been radiocarbon dated between 500 and 1100 A.D. (Vashilian 1978). Probably originating in southwestern Puerto Rico, Ortoiroid culture spread to all four coasts and significant portions of the island's interior, eventually expanding into the Dominican Republic and the Virgin Islands, and exerting a overall sphere of influence as far south as Trinidad (Bullen and Bullen 1974: 4). The Ortoiroid people were agriculturalists who also consumed substantial quantities of shellfish. Indeed, this cultural phase of Caribbean prehistory was originally labeled Shell Culture by Barry (1948).

The term Ortoiroid itself is Spanish for oyster. The type site of the series is at Punta Ortoiro, Cabo Rojo, virtually a small peninsula made up of shell debris. In addition to shellfish and root crops, Ortoiroid subsistence included maize, fruits, nuts, fish, lobster, sea urchins, possibly scorpions, wrenworts, manioc, turtles, tortoises, quarts, small land mammals, etc. It is clear from the archaeological record that these people exploited the sea, the shores, rivers, mangroves, and inland resources.

Compared with pottery as well as later prehistoric ceramic traditions in Puerto Rico, Ortoiroid pottery is rather plain, more functional than artistic. Diagnostic traits include deep banding, diagonal ribs, angular bases, red slip, winged edges flaring outwards, vessel walls averaging 3-6 mm (3 in.) thick, bead bags, necks, and bowl-shaped

TABLE 2
Description of Prohibition Sites Used in the Study

Site and Count	Relation to Sea Level	Date B.P.	ESM
1. Isla Verde (3)	Submerged 0.5 m	1200-600	200
2. Bolsones Isla Verde (2)	Shoreline beach	1200-600	-10*
3. Mar de la Cruz (3)	Prograded seawater	2000*	+500
4. Punta Bolsonchero (2)	Intertidal	1445-1245*	0
5. Aguas Frías (3)	Near shoreline	1200-600	-10
6. Colones de San Juan (2)	Total intertidal barials	1200-600	15
7. Escamela Honda (2)	On salt	1200-700	-20*
8. Playa Blanca (2)	On salt	"	+20*
9. Petroplylas Caca N (2)	Intertidal	1100-700	-20*
10. Petroplylas Caca S (2)	Partly submerged tidal	1100-700	-20*
11. Cayo Montego (2)	Islet	1000-600	-80
12. Joles (2)	Island middle	2000*	+1.2*
13. Cayo Colón (2)	Partly submerged-salt	2245-2275*	-204
14. Cayo (2)	Intertidal	500*	-15
15. Bay of Montego (2)	Islet	1200-600	-208
16. Papaya (2)	Island middle	2000*	+300
17. Mayaguez (2)	Islet	1000-600	-108
18. Cayo (2W)	Island middle	2000*	+15
19. Nequeles (2W)	Slightly prograded	1500-600	+50
20. Ocasio (2W)	Slightly prograded	1545-600*	+10
21. Cayo Razonar (2W)	Submerged 0.5-1.0 m	1400-600	-22*
22. Joyuda (2W)	Intertidal	1200-700	0
23. Petroplylas, Escamela (2W)	Intertidal	"	-10
24. Mar Chapala (2W)	Intertidal	1200-600	0
25. Tortuguero (2W)	Intertidal	1700-600	0
26. Petroplylas, Manzanil (2W)	Intertidal	2000-700	-10
27. Playa Caca Gorda (2W)	Intertidal	1200-600	-10
28. Punta Manzanil (2W)	Intertidal	1200-600	-10
29. Punta Caca (2W)	Intertidal	1200-600	0

ESM = Estimated Shoreline Migration

* Radiocarbon Date

For site locations, see Figure 3

month (Knap 1940: 15; Jones 1964: 508, 1968: 124). Thick clay
paddles for cooking various sea conceptions, clear evidence that these
people were agriculturalists.

Otagoa culture demonstrated the rise of the protohistoric Tapa
dialects. It is at the Otagean phase of development that the first
ceremonial plazas or *toiyas* were built. In Puerto Rico, ceremonial plazas
have been found in Utuado, Lajas, Adjuntas, Barranquitas, Caguas,
Jayuya, Ponce, Cabo Rojo, and Mona Island. Alegre (1951) and Fernández
Isidoro (1976: 48) suggest that these plazas represent a direct link with
Caribbean peoples or Mesoamericans, as the Aztecans plazas were used
for playing a ball game closely resembling the Maya *Pit in pit* and the
Haroan *Fluctu*. Yule Haggels (1872) sees a stronger influence from
South America, where the Olmecs and other aboriginal societies of
Yucatan played a ball game as well. In its many variants, the ball game
was played by a wide assortment of New World cultures, including the
Hohokam of the North American southwest (Lape 1976: 144).

The cult of three pygid stones or *shel* marks another Otagean
culture, but it was at this phase that this megalithic cult expanded
immediately throughout the Lesser Antilles (Hallen and Hallen 1976: 4).
Symbols of power or supernatural entities, these *Urotopians*
gradually evolved into highly sophisticated Tapa sculptures of animals,
gods, and geometric designs.

Whatever its nature and limitations, the influence that
Otagean culture exerted throughout the islands reported political strength
and considerable population density. In a very plain, as yet poorly-understood
form of internal trade (Otagean culture) and external influence (Maya

colours), the Tene chambers of the Greater Antilles came into being. We shall now look at selected coastal sites in some detail.

Isla Verde

Isla Verde is a narrow site submerged off Punta el Morco, 4.8 km east of San Juan (North Coast-Caribbean Morayges). The site is located in shallow water between an offshore reef and Puerto Rico's north coast. Following a preliminary survey in 1978, underwater test excavations were conducted by the author in 1980 (Vega 1987). The site has been subsequently revisited for additional observations up to the present.

For the 1980 field work, two datum points were established: Point A on the southeast corner of a small islet at the end of the offshore reef, and Point B on the north-western tip of a small islet on the coast. A segmented base line was run between datums, and soundings taken with a slath rod at 2 m intervals. The pointed a maximum water depth of 2.4 m and an average depth of 2.4 m. Five 2x2 m test pits were excavated underwater with a water dredge, followed by a smaller 0.6x0.6 m test pit excavated by shovel on the islet.

The underwater excavations clearly revealed a compact midden submerged in place. Located at the median point of a compound basalt embayment, the site was submerged as the embayment expanded (Vega 1987: 24-48). During transgression, the midden was protected from direct surf by both the offshore reef and the islet.

The 1980 excavations yielded Obsidian and a few Elmered pottery shards, polished stone tools made of diabase and volcanic breccia, shell pebbles and grays, and human bones. The were recent Elmered pottery

shards were recovered from the surface of the sea-floor but not below suggesting some stratigraphic coherence. One archeozoogeographic motif was found: a small, flat lake in reddish-clay. This lake is similar to those carved in continental stone cells of the later Tlaxtec culture (Hasson 1984: 208, Fig. 18). A total of eleven massive gastropod or bivalve shells were found, indicators of agriculture. The diet also included manatee (*Trichechus manatus*), sea turtle (*Chelonia mydas*), fresh water turtle (*Chelydron*), porcupinefish (*Diodontidae*), parrotfish (*Micropogonias*), and of course shellfish. Common marine shellfish included *Cantharus tuberosus* Linnae, *Struthosaurus pygma* L., and *Catharus pygma* L. Common bivalves included *Anadonta* s.l. Linnae, *Arca*, *Zebrus*, *Swinsonia*, *Phacusa*, *peruviana* Gmelin, *Trachycardium argenteolum* Smith, and *Lacuna peruviana* L.

Reconstruction of the paleoenvironment indicated that the inlet was part of the prehistoric shore, the real point of a smaller lagoon embayment. As the embayment expanded, the inlet was exaggerated and the lagoon submerged. With the sea to the north and a larger San José Lagoon to the south, the village was gradually surrounded by water. In modern times, at the eastern end of the embayment, Tercera Lagoon was opened to the sea, killing much of the area's once famous coral gardens.

Belisario Isla Verde

This is a small detached site at Isla Verde Beach, approximately 500 m (1640 ft.) S-SW of Barra de Cangujeos (North Ocuil-Cordoba). The site is mostly covered by the road pavement of the public beach, but sparse pottery shards may be seen at the beach's edge adjacent to the

permanence. The site appears to be partly eroded, but during inspection off the beach revealed no ceramics or shell discard.

María de la Cruz

María de la Cruz is a multi-component site located east of the Loma River, Rio Grande (North Coast Loma). In 1948, Alegría conducted test excavations there, finding a preceramic level beneath a ceramic level of Curves white-on-red pottery (Alegría et al. 1955: 154). Animal bones, many of them burned, were present in great quantities. Tools from the preceramic level included haemulid bones, shell awipers, pebble graters, choppers, etc. In 1948, additional excavations by Alegría and Holmsted yielded two levels in poor condition (Alegría et al. 1955: 117). That same decade, Kroe (1959: 115) used the site to date local beach ridges in order to estimate the rate of beach advance near the Loma River outlet. For Puerto Rico, if not the entire Caribbean, this was the first geomorphological study in which archaeological sites were used as markers of shoreward migration. The preceramic/prehistoric importance of this María de la Cruz Site are discussed in Chapter 5.

Punta Embarradero, Monserrate

The site of Punta Embarradero is located in Barrio Monserrate (North Coast Lapaquí). The site, also known as Lapaquí or Monserrate, is on a modified near-a small stream at the east end of Lapaquí Beach, facing a shallow reef that runs east all the way to Costa Azul. The shallow entrance (with Bahadad and Olaned currents between), and is approximately 300 m

1044 ft ± E-W by 200 m (1040 ft ± N-S). The original site consisted of five mounds, one of these has been partly eroded by the sea. Three of the mounds were excavated by Kenney (1945), who retrieved a relatively large number of complete or nearly complete ceramic vessels. At the Shell or Calabozo level, both *crusta* and *red-slip* ware were found, with the red paint generally applied in a row band or in curvilinear patterns on the inside.

Numerous burials were found at Embarradero, "so close together that one skeleton could be distinguished only with difficulty from another" (Kenney, 1945: 70). This prehistoric cemetery included burial sites for small children. In Mound A, towards the center of the site and just the eroded one, several skeletons were found below the refuse, half-submerged in water, a clear indication of a relative rise in sea level. Some of these burials were "fairly well preserved" (Kenney, 1945: 78). Common shells at the site included *Sitostoma* sp., *Tellina* sp., and *Chama* sp. Animal bones were abundant, including manatee, tortoise, fish, and bird.

Agua Fria Lagoon

Agua Fria Lagoon is located east of Bahía los Cabanos (also known as Estero de Yagual) near Santa Rosa Beach (North Coast/Papardo). Surrounded in mangrove, the lagoon has a narrow outlet to the northwest, behind Cabana Chiquita headland. To the east, behind a secondary headland, the lagoon has nearly formed a second outlet. A mudbar is located on the north side of this new outlet.

The edge of this Calabozo mudbar is approximately 4 m (13 ft) from the tidal zone. Surface collecting yielded the standard *crusta* and smooth ceramic ware. A dive inspection of Bahía los Cabanos was conducted, using

both sides and extending past. With the exception of one Osteonod shard in very shallow water, at the eastern end of the bay, no archaeological materials were found under water.

Cabanas de San Juan

Cabanas de San Juan is a narrow peninsula that forms the northeastern tip of the island (East Coast-Pigadia). The site is located northeast of Laguna Grande, near the fishing village and marina of Las Cochas. The site is an Osteonod burial ground, situated in a small valley facing the nearby island of Isosue. A long beachrock pavement runs along the beach. In it are the tilted remains of at least five burials aligned along the beach (Figures 18 and 19).

As the sand became beachrock, the skeletons were for the most part dismembered, but preserved in their original place. In one case, the fetal position is clear. All five bodies appear to have been accommodated inside large, thick walled, unpainted ceramic vessels. Pottery shards from these vessels are embedded in the beachrock pavement. Originally, two skulls were visible on the surface of the pavement, but were shielded away by a jet horizon. As a note on toponymy, these skulls may have been the "cabanas" or heads of San Juan, although the name may also have arisen from the shape of the peninsula itself. Preliminary surface collecting by the author yielded two osseous gratulas, one perforated run chord, and various craftware and stoneware shards. A small stibium is found adjacent to the burial ground. The area of Cabanas de San Juan is now a restricted natural preserve under the Polivocuous de Preservacion de Puerto Rico.



Figure 10. Labeled professional burial 1, Cabezas de San Juan.

5

Encarnación Honda

This site is located on a small island on Cule de los Indios Islet, Encarnación Honda Harbor (East Coast/Cuba). Encarnación Honda and surrounding areas are within the U.S. Naval Station of Roosevelt Roads. The site is located on the north side of the mouth of the islet, between a

mangrove swamp and the sea. This is a terraced midden containing primarily Shesha pottery, with some Cuernavaca and Tapanatepec sherds (House 1962b: 148). House excavations at Xucumaná Honda yielded numerous massive profile sherds: a clay ligura, clay disks, a bone pick, a bone spatula, shell tools and a shell, a shell bowl, a shell pendant, hammerstones, etc. The midden was composed of massive bivalves and caracoles, as well as spolia, and bones of manatee, fish, bird, sea turtle, and fish (House 1962b: 148).



Figure 71. Interred prehistoric burial 2, Culucan de San Juan

Playa Blanca

Playa Blanca is one of three middens on the north side of Caba de los Indios at Ensenada Honda (East Coast-Cuba). The island is separated from the sea by a mangrove swamp. It is a small site, approximately 10 m (33 ft.) in diameter. Only one isolated pottery shard was found at the surface, otherwise, it appears to be a prehistoric midden. Relatively few artifacts were located, including shell tips, a stone step, blunted clam shells and coral fragments. The midden is composed of marine bivalves and nautilus, corals, crab, fish, mangrove, and latex (Rusert 1962b: 308).

Ensenada-Caba de los Indios North

These petroglyphs are located on the north side of Caba de los Indios site, Ensenada Honda (E-Cuba N.E.). First reported by Alphonse Fournier in the 1880's, the petroglyphs consist of two anthropomorphic figures and part of a third, at the edge of the water. (Fournier et al 1884: 312). The petroglyphs are carved on large boulders at the edge of the mangrove swamp, across the reef from the site of Ensenada Honda.

Ensenada-Caba de los Indios South

These petroglyphs are located just off the shore of a small island near Ensenada Honda site (East Coast-Cuba). They are approximately 400 m (248 miles) south of the first set of petroglyphs. An anthropomorphic face with six lines radiating from the chin was reported by Rusert (1962b: 348).

More recently, a second human figure was discovered nearby, as reported by Woods (in Trousdale et al., 1984: 3-14).

Cape Santiago

Cape Santiago is a small island off Humacao Beach (East Coast-Humacao). A circular shell midden is located on the low, northern part of the island. The midden is 30 m (100 ft.) in diameter and 35 cm (15 ft.) deep. Contents are mostly Santa Elena and Espinoza (Bosse 1993: 391). The island has a maximum elevation of 34.9 m (114 ft.), and is separated from the parent island of Puerto Rico by a channel roughly 400 m (1312 ft.) wide and 1.8-3.4 m (5-8 ft.) deep. The bathymetry of the area clearly indicates that Cape Santiago was connected to the parent island in prehistoric times, the rocky point of a now submerged beach. Nearby Humacao Beach is an area of sparse woods (Merrill 1879: 46). At present, Cape Santiago is used by the School of Tropical Medicine of the University of Puerto Rico, as a laboratory for primate research.

Jales

The site of Jales consists of 8 middens, 1.3 km (0.8 miles) inland from the Bay of Jales (South Coast-Ogajsona). The middens are spaced an average of 100 m (328 ft.) over the coastal plain. One of the middens contains prehistoric remains. One brick and three Spanish pottery sherds were also found. The rest of the middens appear to be prehistoric. Bosse (1983a: 336) has written: "Jales is of obvious significance. It is situated too far from the shore to be a place where shells were gathered." Yet, he adds

that "the shells are well bleached as in Indian sites, and the heap has an appearance of antiquity." A tentative prehistoric origin was assigned by Rouse to some of the middens. Three decades later, Jeline remains a problematic site (Jeline 1985: 36).

Cayo Culcasi

Cayo Culcasi is a prehistoric shell midden at Jeline Bay (South Coast, Jeline). The site itself forms an islet at the south end of the Mar Negro (Black Sea) mangrove lagoon. The midden is roughly 58 X 40 m. (184 X 131 ft.). Part of the shell midden is now submerged.

Cayo Culcasi was excavated in 1973-74 (Wilke et al. 1979), yielding human remains, blades and microblades, choppers, scrapers, ground and flaked shell pebbles, and a variety of shell tools. The people of Cayo Culcasi ate turtles (*Chelonia mydas*), fish, snail, bird, and lotus (*Potamogeton*). The fish included barracuda (*Sphyraena*) (sp. *dentocarinata*?), and jack (*Carangidae*). The most common shellfish at the site were *Arca* sp., *Octarius* juva, *Strombus* giga, and *Conus* tuberosus. The remains of three non-human bats were also found, suggesting simple disposal without honorary rites of any sort (Wilke et al. 1979: 89).

Legend has it that the Puerto Rican pirate Roberto Culcasi buried a treasure here. This has led to considerable destruction of the midden by treasure hunters in search of imaginary gold.

Cayote

Punta Cayote is located at the west end of Santa Isabel Beach, near the Camino River (South Coast Santa Isabel). The large midden at Cayote is characterized by San Chico ceramics (Roney 1940: 112), which are roughly contemporary with the late Ocotacoatl. Additionally, Elmerol, Guisanoil, Cayot and Esperanza pottery shards are present in lesser quantities.

Most of the site is now covered by houses. On the south side of the point, the sea has cut into the midden. Artifacts in the area included ceramic griddles, a clay dish, stone splinter, stone ball, bone disk, bone peg, bladed clam shells, steel fragments, etc. Faunal remains included marine bivalves and unidentifiable bird, crab, fish, turtle remains, and one turtle (Roney 1940b: 832).

Bay of Muxtelon

The Bay of Muxtelon is located east of La Pazguera, off the Porton peninsula on the south coast (Lagun). An Ocotacoatl ceramic midden has been reported for one of the twenty-plus mangrove islets in the bay (Mason 1941: 271). The midden is composed primarily of unidentifiable, with small quantities of oysters, clams, mussels, and garden hives. According to the topographic map of the State Natural Preservation Office in San Juan, the site is located on the largest of three islets off the old gulch called south of the salina. Locals interviewed by the author call this islet Cayote, but reported no knowledge of a shell midden there or in the two islets nearby. Scurfing around Cayote and then traversing it, the author was unable to locate the site. According to Mason (1941: 271), the islet is approximately

118 m (130 yards) long and completely covered with shell, and is probably one of the offshore, mangrove-free islands of the Bay.

Estero

Papaya is located between the Bay of Maricao and La Parguera South Coast-Lago. The site includes 18 mounds on "slight rises of land at the edge of an extensive mud flat" (Beane 1955b: 248). No prehistoric ceramics were found, but an iron belt and one European pottery shard were found on the surface of a mound. Presuming that these two items were intruders, Beane has suggested a possible prehistoric origin for these mounds. Tools include stone chips, a stone disk, fluted clay shells, shell tips, and two coral fragments.

Manatí

Manatí is a small, oval shaped islet near the town of La Parguera South Coast-Lago. The Field Laboratory of the Institute of Marine Biology at the University of Puerto Rico is located there. The islet is separated from Puerto Rico by a very shallow canal roughly 100 m (330 ft) wide. The south and east coast of Manatí are mangrove, as is most of the coast of La Parguera and the south coast in general. The highest point on the island is 27 m above MSL.

A shell midden is located on the southeast corner of the islet. The midden has very few potsherds, which may be classified as early Colonial. Most artifacts are made of shell, including coils, cups, polished pendants, disks, and shell-pipelets (Cowan 1955: 20). Common individuals at the

also include *Strombus costatus* Gmelin, *Strombus* sp.?, *Lorai*, *Strombus papillat* L., *Chamaea brevifrons* Lamarck, and *Conus*. Johnson L. Conus behavior include *Anadara nitidula* Hilleb, *Chama macrophylla* G., *Chama senaria* Boudier, *Arca nuda* Swenson, and *Chama granulata* G. It is worth noting that *Anadara nitidula*, while common at Manayoa, is now considered rare or absent in Puerto Rico (Cousins 1963: 73).

CORRAL

Corral is a prehistoric midden near the town of Boqueron (Southwest Coast-Cabo Rojo). Located 1 km (0.6 miles) north of Boca Bay and 1 km from the west coast, the site is a single midden of both levels and massive shellfish. The midden is 20 m (65.6 ft) in diameter and 25 cm (1 ft) deep. *Anadara* includes *hemisphaerica*, *Macta*, a shell cup made of *Strombus*, shell points, and a coral fragment. No water supply is available (Rouse 1964: 375; Davis 1963: 23).

BOQUERON

Boqueron is a large prehistoric site near the bay of the same name (West Coast-Cabo Rojo). The site is now 204 m (669 ft) in diameter and 58 m (194 ft) from shore. Rouse (1963: 374-6) reports *Chamaea* and *Conus* were, as well as clay griddles, a discoidal clay stamp, an anthropomorphic stone pendant, a shell *Macta*, a shell cup, etc. An *Orconoid* child's burial was excavated. This was in very poor condition and had no associated artifacts.

Cotacotán

The site is located at the top of Punta Cotacotán (West Coast-Cabo Negro). Enclosed by Lichup, de Huerco, Spínola, Montalvo, Basso, and numerous pit features, the original midden had a diameter of 100 m (328 ft) and consisted of six mounds, five of these forming a horseshoe (Roney 1960: 113, Basso 1966a: 165).

The site is situated on a low, sandy point between the ocean and a mangrove swamp. Excavations by Basso and others yielded a wide assortment of pottery forms, including bowls, cooking griddles, cylindrical discoidal stamps, a spherical clay head and a clay three-pronged or comb. The site has also yielded stone celts and chisels, stone polishers, stone cylinders, shell tip hammers, a shell disk, shell celts, etc. (Basso 1966: 167).

Recent observations by the author revealed a curious geomorphological feature unreported by previous researchers. The south or seaward end of the site is beachrock, with numerous pottery shards embedded in it. Yet, the actual shoreline is not there, but some 12 m (39 ft) south of the beachrock pavement. This suggests two shoreline movements since the time of prehistoric occupation. First, the sea advanced to the edge of the midden, where the beachrock pavement was formed, then the shore prograded, leaving the pavement inland. The implications of these observations are discussed in Chapter 5.

Cerro Batatas

Cerro Batatas is a small islet located off Jacó (East Coast-Cabo Negro). Also called Puma and Puma in historic times (Coville 1895: 35),

Cayo Batones is roughly 122 m (400 ft) E-W by 62 m (200 ft) N-S. The site is presently at a distance of 427 m (1400 ft.) from the parent island of Punta Iles, separated by a protected channel with a maximum depth of 4.6 m (15 ft.)

The west or offshore side of the site is protected by remnants of mangrove. The center has sufficient sand to support a small wooded area of palm and other coastal vegetation. The rest of the islet is sand. Immediately south of Cayo Batones, at a depth of 2 m (6.5 ft) and shallower, Olmeca pottery shards may be seen on the sandy bottom. The surface material is uncovered on a seasonal basis, which makes it difficult to estimate the extent of the deposit, but apparently covers at least 30 m (98 ft.) in circumference.

Jayula

The Olmeca site is just off the town of Jayula, (West Coast Cuba Bay). In 1913, Leidy excavated at Jayula, noting that "the sea is rising into it and many objects have been found on the beach. A few years ago several skeletons were dug up which were reburied in the Cuba Bay cemetery" (in Bruce 1952a: 300).

Jayula was also excavated by de Beary in 1916 and de Harter in 1917. Based on her own surface collecting and on examination of materials from previous excavations, Bruce (1952a: 300) reports that most ceramics at Jayula are Olmeca, with some representation of Coclé shards and a few Elmina shards as well. The material included narrow griffles, dentoidal clay rings, a cylindrical clay bowl, stone paddles, a stone axe and celt, a cylindrical stone bowl, an ellipsomorphous stone pendant, shell coils, a shell

lead, etc. Based on ceramic analysis, Brown (1992a: 240) has suggested that *Lapacha* may be somewhat later than the type site of Otomani to the south.

Interoptima, Punta Barrosa

These petroglyphs were located at the tidal zone in Punta Barrosa, (West-Coast Brazil). Carved on beachrocks, the petroglyphs appear to be of Ceramic Age origin, although their exact cultural affiliation is not entirely clear. As nearby sites, there is evidence of Ceramic Age occupations from early Saladoid through Otomani and Chiriqui stages (Rodriguez 1998: 4). Moreover, at the site of Estancia, early 16th century Spanish artifacts have been found, including coins, indicating changed ethnicity into historic times (Carlson 1998: 12).

Mar Chiquita

Mar Chiquita is a lunate moonshaped bay at Maratí (North Coast, Maratí). An Otomani midden is located approximately 400 m (1312 ft.) east of the beach, directly south of a small islet. This islet is a remnant of the same andesite ridge that forms the Mar Chiquita beach. The midden is partly exposed at the beach scarp, and at the forefoot of rock cut-off. Numerous pottery shards are embedded in the long, narrow beachrock pavement fringing the beach.

Tortugues

This Guatemalan site is located NW of Tortuguero Lagoon, between Mar Chiquita and Puerto del Tortuguero (North Coast, Mexico). A narrow beach fronts short, limestone cliffs with minor karst formations. It is clear that this site had a number of prehistoric centers at the time of prehistoric occupation. A preliminary impression suggests that some of these centers were occupied by the aboriginal population. Eventually, the centers collapsed due to marine erosion. Abundant pottery sherds are visible at the base of the cliffs.

Manatí

Manatí is one of various prehistoric sites between the volcanic ridge of Puerto Puerto Nuevo and the wetlands of Boca del Caimán (North Coast, Yucatán Region). Excavations at Manatí have been recently undertaken by the Centro de Investigaciones Indígenas de Puerto San Diego (1988). As is usually the case, the jet hunters have also done a considerable amount of digging at this important site.

The site of Manatí is not from the beach by a paved road. Pottery sherds are visible by the side of lanes along the road. On the beach, the beachrock pavement is decorated with a variety of petroglyphs. Seven petroglyphs were identified by the author, presumably, others may be found beneath the sand. The exposed petroglyphs are located by the sea at high tide.

Due to the fragile nature of beachrock, and the relatively smooth cut of the petroglyphs, it is likely that the prehistoric artist worked during the

earliest stages of the process. At this stage, beachrock has the consistency of semi-dry cement, and the artistic technique may be an intermediate stage between petroglyph etching and petroglyph carving.

A dive search was conducted in the reef-protected embayment between the beachrock coast of Manabí and the volcanic ridge of Punta Santa Rosa. Only two shards were found: one barbed, the other a microfluted, Colossal-like shard. This sample alone suggests, however, that the embayment is not as suitable for canoe as a surface view might suggest. Breakers rise close to the shore, and in rough weather the waves explode against the volcanic ridge. If canoes were used here, they were small. It is far more likely that the prehistoric inhabitants had access to the sea by Punta Santa Rosa and the mouth of the Cabaes River to the east. Once they knew the coast, the prehistoric navigators had no difficulty launching or landing a canoe at the Boca del Cabaes.

El Estero Grande

El Estero Grande is located east of the Cabaes River (North Coast Yuga Bay). A large number of Colossal pottery shards are embedded in beachrock for more than 500 m (1640 ft.) along the beach. The wooden vessel is located on private land near the border of the Yuga Bay Yuga Alto Municipio. No ceramics were found underwater.

Punta Manayes

Punta Manayes is located north of Grande Airport (North Coast Road). The site reaches to the top of Punta Manayes, and is partly eroded, to the south, it is partly covered by landfill of a nearby hotel. A narrow beachfront pavement runs along the point, with Oakenood pottery and shells embedded in it. Isolated pottery shards may be found under water.

Punta Cerezo

Punta Cerezo is an extensive point that protects the mouth of the Coastal River (North Coast Two Bay). The site is located on both sides of the mouth of the river, with most of the refuse on the west bank, near the tip of Punta Cerezo. The Coastal River is a secondary outlet of the larger Plata River to the west. Immediately south of the site, the Coastal widens as it drains a relatively large watershed. The site is an Oakenood midden extending at least 40 by 40 m (130 ft.). The eroded walls of the midden are clearly visible on the west bank of the Coastal River. A shore search was conducted off the site. Due to the presence of the river, underwater visibility tends to be poor, but the high tide allowed us sufficient visibility to verify that the midden-edge is the total zone.

Historic Sites and Documents

Some caution is to be taken in inferring that there are any historic sites simply because to have a historic site is by definition an inferred concept. This is put in perspective as follows in Latham:

Richard C. Flannery (1979: 28)

Non in causa mandatis
 Quia non modo de legem
 Causa non mandatis,
 Non antiquitate de mand
 Saltem Falsi dicitur in Causa. 1081. 38

Historic evidence of coastal change includes archaeological sites as well as cartographic, documentary, and photographic data. In theory, modern coastal change is intrinsically historic, and need not be placed in a separate information domain. Whether a coastal structure is Puerto Rican 1900 or Colonial Spanish 1600, in either case it pertains to a historic (historic) culture and as such may be dated with great precision.

This section examines diverse historic markers of coastal change in Puerto Rico (Figure 12). These sites provide further evidence of coastal

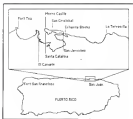


Figure 12. Location of historic sites used for the study

change up to the present, adding support to the Model of Caribbean Submergence presented in Chapter 5. Moreover, humans were not important in their own right, and Cocoloba should be included here, in order to promote awareness of sea levels in historical archaeology (Table 3). We shall now look at selected historic evidence of coastal change for the main island of Puerto Rico.

Fort Tio

The Bay is a large, lacustrine embayment west of San Juan Harbor. When the city was founded, the Spanish authorities at San Juan believed that this bay was completely enclosed from the sea by a chain of reefs. In 1647, however, marine surveys found an entry large enough for warships to get in, an obvious threat to the defense of San Juan (Juana de Guerra 1983, Lopez 1973: 218).

The gap in the reef was too wide and deep to be closed with rubble, compelling the governor to build a fort (Novoa 1890). Started in 1652 and completed in 1664, Fort Tio was built as an *islet 'en medio de la mar'* – in the middle of the sea – near Punta Ballenas. It was a square structure made of masonry, large enough to hold 300 soldiers, and provided with an iron cannon and a bronze piece (Pons de Guzmán 1984a). A quay, 3.5 m (11.5 ft) long, was built of stones dumped within timberwork. The Bay is depicted as a 1660 map produced by Cayetano Dantis (Figure 12). This rather impressionistic map was made two years before the construction of the fort was initiated, but from the archival data it is clear that the fort was built at the ‘Tio Islet’ depicted by the letter ‘H’ on Dantis’s map.

TABLE 3
Description of Historic Sites (Data on the South)

Site	Coast	Sea Level History	Date
Fort Teo	N	Caisson fort built on reef Presumably submerged	1884
San Juan de la Cruz El Cabaño	N	Masonry fort on reef Sunk under MSL in the 1830's Now used to fill through sand fill	1884 ±
La Fortaleza- Santa Catalina	N	Masonry seawall undermined by sea in the 1750's Modern sand fill and repair	1834
Morro Castle	N	Masonry seawall undermined by the sea throughout history Modern sand fill and breakwater	1869
San Cristóbal	N	Masonry seawall. Extensive damage reported in 1831 and 1764 Profile of 1769 shows lower MSL	1823
San Jerónimo	N	Caisson sand brick fort Base is 7 ft on water table	1887
Indiano Rocks	N	Placed by military agents to block entry to lagoon Sea level caps outer MSL	1797 ±
La Torreola	N	Fortified watchtower Left presumably submerged	1888 ±
Fort San Francisco	W	Masonry fort hit by earthquake and tidal wave in 1918 Outer wall within tidal range	late 18th century

In 1797, the year of the British invasion under General Ralph Abercromby, the Spanish built yet another fort, this one on the main island east of Punta Salinas, since then known as Isla Sotavento (Cardona 1988: 17). Made of masonry, its foundations were 22 ft (7 ft 6 in) above sea level (MHL 1990). The ruins of this small fort are still standing, and should not be confused with the older Tin Fort. In a letter to the King of Spain, Gov. Pascual O'Quin (1884a) indicated that two bridges were used to transport the construction materials to the island, as the surrounding water was too shallow and rocky to allow large boats. Both bridges had a span of 487 m (1600 ft.), meaning that the fort was roughly 915 m (3000 ft.) from shore.

Obviously, the ruins of Fort Tin are beyond Isla Sotavento, which is less than 100 m from Punta Salinas. A 1942 aerial shot of Puerto Rico (General Library, University of Puerto Rico, Box Pindora Campus), clearly depicts the two reefs between Punta Salinas and Cabana Island. Presently, the submerged remains of Fort Tin are to be found in the shallow of the western reef off Punta Salinas.

Currently under the protection of the U.S. Armed Forces, Punta Salinas is a beach with signs of considerable erosion. Nearby Laviterna Beach is also being eroded. The problem appears to be largely due to the Isla de Cabana causeway, at the west end of San Juan Harbor. Built during World War II (Cardona 1988: 26), this causeway has radically changed the current and wave patterns in the Tin harbor embayment, resulting in accelerated erosion (Morselock 1978: 26). An additional factor to be considered is the reflection of the Tin River in late historic times, which now drains east of the embayment, greatly reducing the sediment supply to Tin Bay.



Figure 16. Plan of San Juan, depicting La Tierra Alta (A), El Encanto (B),
 Monte Grande (C), San Juan (D), and San Juan (E).
 (Source: Archivo General de Indias, Mexico y Plazas, 1614-1615)

San Juan de la Cruz o El Castaño

San Juan de la Cruz, commonly known as El Castaño, is a small fort built across the harbour from Stone Castle. Originally located on a small islet behind Colman Island, El Castaño was built of wood in 1600 on shorelands, burned by the Dutch in 1633, and rebuilt in masonry in the 1660s. Both the islet and the fort are clearly depicted on Luis Venegas de Osma's 1679 plan of San Juan Harbor (Figure 14). In a letter to the King of Spain, Gerv. de Medina (1685) indicated that one of the walls of El Castaño had been badly damaged by the continuous attack of the sea. No mention is



Figure 14. Plan of San Juan Harbor, depicting Port El Castaño as a small islet. By Engr. Luis Venegas de Osma, 1679 (Archivo General de Indias, Mapas y Planos 802-74)

made of hurricanes or storms. Don de Medina indicated that his predecessor, Don Gaspar de Arredondo, built a stone groin around El Castaño, as protection against the sea's attack on the fort. It is clear that the small islet of El Castaño was being ground at a fairly rapid rate in the late 17th century.

Originally, El Castaño islet was 487 m (1598 ft) from shore, as indicated by the span of the wooden bridge built to transport the construction materials for the fort. (Perez de Guzman 1664) This is confirmed by various maps of San Juan produced in the 16th, 17th and 18th centuries. The most recent map in which the islet of El Castaño is shown, was produced in 1843, by Sergeant Charles Lapinski. Archivo General de Puerto Rico, Colecciones Particulares-Bufiles. Papeles 71. That same year or soon after, the U.S. Army Corps of Engineers filled the challenge surrounding El Castaño, in order to facilitate access to military installations at Cabana Island (Caribbean 1958: 161).

In a plan and two profiles of El Castaño produced in 1858 (Archivo General de Puerto Rico, Colecciones Particulares-Bufiles. Papeles 71, Capt. Henry B. Morganot, indicated that some of the walls are cracked, which he attributes to "subsidence or the settling of the foundations." And he adds that "the waves have eaten into the base of the fort around its entire perimeter." The profiles clearly indicate that, by 1858, the foundations of El Castaño were under water.

La Fortaleza, San Juan Harbor

Completed in 1548, La Fortaleza is an early stronghold overlooking the inside of San Juan Harbor. In 1884 or soon after, Eng-

Don Sebastián Alonso built the massive walls in front of La Portalesa, which came to be known as Santa Catalina. This was part of an ambitious undertaking to enclose the city of San Juan with fortified walls and bastions, a project that was finished by 1638 (Vila 1974: 174-181). In a 1752 letter to the king of Spain, Don Caldeira indicated that the walls at Santa Catalina had been undermined by the sea, and were in dire need of repairs (Caldeira 1752).

Morro Castle

The fortification of San Juan's Morro-or headland began in 1540 with a round tower overlooking the narrow entrance to the harbor (Rios 1973: 70). Construction continued until the late 16th century. The oldest walls date from 1608, while most of the superstructures as we see it now date from the late 18th century. Fear of attack as well as actual assaults persuaded much of Puerto Rico's early governors to add new elements or repairs to the stronghold. San Juan was attacked in 1588 by Sir Francis Drake, in 1597 by the Earl of Cumberland, in 1625 by the Dutch under Hendryck Hendrickson, and in 1747 by Gen. Ralph Abercromby. In 1898, the first shot of the Spanish-American War was fired from Morro Castle. Finally, the U.S. invasion troops were landed at the much weaker part of Old San Juan, on the south coast (Rios 1974).

The sea has been a less responsive but constant adversary against Morro Castle. The fortress has atop a large headland of volcanic origin, of which four separate geological formations have been identified by Keys (1989b: 184). Throughout its history, there are numerous instances in which El Morro's seawalls had to be repaired against the attack of

the sea (Lopez Castro 1975: 211). In recent years, from 1977 to 1988, some \$33 million have been spent on repairs at El Morro. For the most part, damages have been due to wave action against the walls. The problem appears to have increased through the centuries, prompting the construction of a 248 m long (795 ft.) breakwater to protect the scotch walls against northerly seas (Chen 1988: 3). This breakwater is currently under construction.

San Cristóbal

In 1834, San Cristóbal was a small installation at the northeastern end of the San Juan headland, overlooking both the north coast and the city. Larger than El Morro, San Cristóbal was completed by 1763, and played a prominent role in the defense of the city against the Dutch invasion of 1797. In 1850, a report on Puerto Rico's defenses indicated that the scotch of San Cristóbal, built in 1628 by Christopher, was in need of repairs (Lopez Castro 1975: 211).

In 1799, San Cristóbal's scotch was in need of repairs once again. That time, the governor's report was accompanied with three profiles by engineers Juan Francisco Mestre and Tomás O'Daly (Figure 18). Two of these profiles provide mean sea level (MSL)-1740. Preliminary observations suggest a 25 cm increase in sea level in the last 250 years.

San Justino

Fort San Justino is located on the west side of El Neguero, which is the entrance to Culebra Lagoon. This body of water separates San Juan

what from the Sancti-Spargus side. Today, water and resident shiks are seldom aware of these shiks, camouflaged into a single land mass by bridges, land fills and urban development.

Fort San Jerónimo played a key role in the defense against the British invasion of 1761. Although the fort itself dates from the 17th century, its foundations were laid in 1587. The original structure consisted of a platform with two main casemates, but large enough to accommodate six pieces (Alexander de Valdivia 1847). Both the foundations and the masonry consist of volcanic tuffstone, cemented with mortar and reinforced with rectangular bricks.

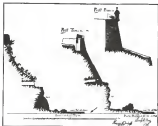


Figure 15. Profile of Fort San Cristóbal, with 1588 sea level. By Engrs. Juan Fontaine-Morales and Tomás O'Daly, 1768 (Archivo General de Indias, Mapas y Planos, 20 387).

Diving around the perimeter of the fort, the author estimated the foundations to be 70 m (230 ft.) under water. Although the techniques for underwater construction are known since Roman times, Spanish documents give no indication that this was the case at San Jerónimo. Off the NE corner of the fort, a brick rubble zone was discovered at a depth of 2.4 m (8 ft.), evidence that the superstructure or the wall itself has been damaged by the sea in historic times. The area's exposure to storm waves is also indicated by a small, fishing pier connected to the south or eastern side of the fort. Solidly built like a wharf, the pier was frequently used by fishermen less than ten years ago. On occasion, small sharks were caught from it. Today, the partly-destroyed pier is some 12 m (39 ft.) under water. An isolated, perforated pottery shard (Cortésoli) was found by diving in the middle of the channel. Presumably, this shard travelled from an industrial or submerged production site in the Concha area.

ElCanoa Blockade (Concha Lagoon)

On the east side of El Boquerón, opposite Fort San Jerónimo, the channel is partly blocked by large volcanic blocks. According to Angel Rivera, a soldier-historian who served as Commanding Officer of Fort San Cristóbal during the Spanish American War, the volcanic blocks were placed there in 1797, to partly close El Boquerón in anticipation of the British invasion under Gen. Abercromby (Rivera 1973: 40).

In 1798, two years after the British siege, a plan to block the entrance to El Boquerón was proposed (Vargas 1798). Unless this meant to further close the channel, the blocks were placed in 1798 rather than in 1797.

Either way, it is clear that the blocks were blasted from the nearby volcanic formation at El Cordón, and placed in the channel in the late 1750s.

In a survey of the sea-level watches or signs that have been found on the blocks, Kays (1894b: 116) concluded that sea level (MSL, 1985) has remained stable since 1757. However, he qualifies his "present sea level" as "plus or minus two feet" (Kays 1894b: 80). From a geological perspective, a 0.6 m (2 ft.) rise in sea level might be considered negligible, but to the archaeologist it may have tremendous significance. Unlike Kays, the author inspected the walls by scuba diving, observing sea level signs at least 10 cm (4 ft.) under the present MSL. However, underwater observations suggest that some of the blocks may have moved in the last two centuries, due to the undermining of coral beneath the blocks. Additionally, the particular coral tends to get fairly rough, which makes it difficult for the scuba diver or snorkeler to conduct accurate observations except at the best weather.

La Torrecilla

La Torrecilla was a fortified watchtower at Boca de Cangrejos, east of San Juan. The purpose of this military installation was to defend the area's plantations against warlike raids by Carib Indians during the late 16th and early 17th centuries.

In a letter to the king of Spain, plantation owner Miguel Pizarro (1684) requested "*una o más fortines blancos de cañón*" – one or two white men of iron – to protect the plantations against Carib attacks. In 1685, the Crown approved Pizarro's petition. Six years later, the Ofendos Baster or Royal Officers of San Juan (1691) prepared a report on the *estado de Puerto Rico* for General Diego Pizarro, including the construction of La

Torreón at Cagayán. Gov. Pizarro arrived from 1698 to 1699 (Vila Vilar 1954: 80), which indicates that La Torreón was built sometime between 1698 and 1699.

In Capt. Vicente Cantón's 1640 map (Figure 1E), La Torreón is depicted under the letter "A" at the extreme left. The legend indicates that the watchtower had "*puerto para barcos* " - a *puerto* pier. In a 1697 map of Puerto Rico in Sancho's *Antares General de Indias* (Mapas y Planos, Santo Domingo 124), the point adjacent to Boca de Cagayán is called *Punta de la Torreón*. In 1775, Miquel (1994) reports that La Torreón is no longer standing. The watchtower's foundations have never been found. Presumably, they are under water. Today, Boca de Cagayán and nearby McDonald Beach are areas of severe erosion (Marschall 1979: 20). An unusually long beachrock pavement runs from Cagayán into the sea for 400 m (1,300 ft). To Kaye (1994: 117), this pavement indicates "a profound change in shore outline with the removal of at least 1,200 feet of shore."

The first bridge at Boca de Cagayán was built of wood in 1690, as reported by Gov. Joseph de Rivera y Moxos (1690). Four years later, the new governor reported that the bridge was destroyed by a storm, and rebuilt that same year in wood (de Valencia 1694). In September of the present year of 1699, surfs from Hurricane Hugo caused sufficient erosion to weaken the foundations of the current bridge at Boca de Cagayán, now to be replaced by a metal one. The historic data clearly indicates the unstable nature of Cagayán, further substantiating the legend that La Torreón's foundations are now submerged, or at best buried under sand at the intertidal zone.

Fort San Francisco

The basic structure of Fort San Francisco is located behind the public school at Sector Tamarindo, Aguadilla, on the west coast. The original installation was built in the late 15th century, to provide protection to the exposed bay area, once a prominent harbor. During a marine survey of adjacent Tamarindo Reef conducted by marine biologist Walter Cardona and the author (Page 1997), it was noticed that the sea now washes against the outside wall of the fort, clearly indicating a change in relative sea level.

Tamarindo and other sectors of Aguadilla subsided due to the earthquake of 1905, followed by a tsunami or tidal wave which wiped out the wooden houses adjacent to the fort. One witness wrote that, a few minutes after the earthquake, "horrific sea waves attacked the beach, tore off the houses, and carried it to the depths of the sea" (as Blevins 1938). The late 17th century also appears to have been a time of seismic adjustments in the Maricao Passage area, with "great earthquakes" reported for the west coast (Cabrillo de San Germán 1690) followed by "constant earthquakes" eight years later (Cabrillo de San Germán 1698).

CHAPTER 5

A MODEL OF CARIBBEAN SUBMERGENCE

Within the last 15,000 years, the dominant pattern of Caribbean coastal geomorphology has been one of submergence. In spite of geologically recent tectonic uplift and prograded coasts in the Caribbean Antipodals, the overall process has been a flooding of coastal zones. Significant variations are recognized in the sea level history of Caribbean islands, particularly between the continental and the oceanic islands (as discussed in Chapter 3). Nonetheless, it is clear that the submerged shelves of all island groups share sufficient oceanographic, geomorphological, paleoecological and prehistoric cultural traits to warrant a regional geoarchaeological model of Caribbean shoreline migration.

The Model of Caribbean Submergence (MCS) is a predictive as well as an interpretive model for the archaeological reconstruction of sea levels and coastal change. In its widest scope, the model is intended for the entire Caribbean archipelago. Its aim is not simply to predict high-probability sites for locating nuclear sites, but more importantly, to present a systematic approach to coastal geoarchaeology in general. In that last regard, the model is by no means intended exclusively for diving archaeologists.

The central premises of HCSH are as follows: 1) that reefs are transient boundaries that fluctuate across continental or island shelves in response to sea level changes, and 2) that reefs were the habitat of many, if not most, Caribbean prehistoric people. Considering that the sea has risen an estimated 50 to 100 m (165-445 ft.) in the last 10,000 years, it follows that vast tracts of the Caribbean prehistoric landscape are now under water, most likely including the richest archaeological record for the region.

The model follows a logical progression from the comparative study of intertidal and shallow water sites, to predicting the location of deep water sites. Beyond method and theory per se, this chapter also explores diverse strategies for promoting archaeological awareness of sea levels and submerged land sites in the Caribbean.

Conceptual Framework

*I came to know these water holes by playing on the seashore, and knowing myself
+ on and then finding a number of shells or a pebble shall than ordinary; whilst
the great mass of shells lay all unobserved before me.
Lewis & Clarke 1843:1137-38 (Rasmussen 1977: 3)*

*On all these shores I have seen signs of past and present life of the few of time
elapsing, yet retaining all that has gone before.
Richard Coates 1968: 144*

Models provide both tools and perspectives with which to examine archaeological sites (Binford 1982: 312). Ideally, an archaeological model should be firmly based on field data, but at the same time transcend its empirical foundations in order to generalize beyond the spatial and temporal constraints of the observed sites.

The required foundations of MCS have been presented in previous chapters. These include general Salween sea level changes (Chapter 2), general Caribbean archaeology and ecology (Chapter 3), and the coastal and marine surveys in Puerto Rico (Chapter 4). Inevitably, modeling implies choosing certain terms and theoretical problems and underplaying others. This first version of MCS is particularly concerned with settlement patterns, shoreline migration, submerged site production, and modification of sites. These are the primary areas that need to be researched in order to have a solid grasp on the problem. Prior to presenting the results of the comparative study, it is essential to review the fundamental concepts behind it.

Site Matrix and Eroded Materials

In high energy coasts, it is possible for artifacts to be transported offshore by up-currents, or down the coast by longshore currents. In some cases, eroded materials may create the impression of intertidal or submergical sites.

Of the thirty-six archaeological sites included in this study, only one could not be properly identified as terms of site matrix. The site in question is Salween Isle Yard, which appears to be either buried under pavement or eroded, with only sparse pottery shards visible at the intertidal zone. The presence of these Oriskany shards on the beach indicates that shoreline migration has taken place, but not much else can be said about this particular site. Was it a small site in the immediate area of the observed pottery shards, or a larger site further up the coast, towards Boca de Cangrejo?

Sea Level, Shoreline Migration, and Erosion

The Model of Caribbean Submergence makes a fundamental distinction between sea level change and shoreline migration. Sea level change is an expression of vertical (up-and-down) movement of the sea in relation to the coast, whereas shoreline migration is an expression of horizontal (prograding or retreating) movement of the shore. In areas of passive erosion, shoreline migration may occur without a rising sea level just as a rising sea level may result in little immediate shoreline migration on a stable coast. On low-lying coasts, the correlation of sea level to shoreline migration may vary from 1:100 to 1:1000. In other words, a rise in sea level of 1 m (3 feet) may drive the shoreline from 50 to 500 m (100 to 1000 ft.) inland (Giblin 1984).

From a geoarchaeological perspective, the Puerto Rico data chosen for the study exhibits an interesting division between historic and prehistoric sites as sea level markers. While prehistoric sites are typically ground-level middens or beachrock petroglyphs suitable for measuring shoreline migration, historic structures are "standing yardsticks" suitable for measuring sea level. The highly resistant nature of these Spanish forts is due to two factors: 1) the massive walls made to withstand cannonballs, and 2) the fact that most of them were built atop volcanic

beach ridges as the single most important factor in the destruction of intertidal archaeological sites. This type of erosion is largely the product of longshore currents, for the most part generated by waves striking the coast at an angle. Waves in turn are controlled by four primary variables: 1) wind direction, 2) wind duration, 3) wind velocity, and 4) fetch, i.e., the distance

increased by waves without obstruction. In open-seas, depth is measured in hundreds or thousands of nautical miles, allowing the wind-generated waves to evolve into full seas. For instance, a 50 knot (mph) wind gale requires 2400 km (1492 nautical miles) to produce a fully developed sea with waves averaging 16.6 m (545 ft) in length, and big waves up to 33 m (109 ft.) and possibly higher (Dunham 1994: 52).

Waves break when the water depth is roughly 1/3 their height, when dissipating into smaller waves which will break in shallower water and so on until reaching a coast, offshore bank or reef. Finally, although waves can direct enormous amounts of energy against a coast, at a depth of half the wave length (horizontal distance between two crests), their energy is reduced to 1/32 of the surface strength (Gardner 1994: 39). For submerged land sites, this implies that the greatest exposure to direct wave energy occurs during transgression, in exposed or partly exposed areas.

Beach Settlement Patterns

Throughout history, coastal dwellers have built their homes as close to the edge of the sea as possible. This phenomenon is so prevalent and universal, that the author is tempted to present it as a cultural law. From fishing villages to elite dwellers, from nomads to resorts to middle-class homes, people will build as close to the water as possible. Weather systems, hurricanes, erosion, or tidal waves will change their minds. The destruction of a Port Royal, Haida, or Plover always yields a novel lesson — within a geological age. Even today, with the warning of geologists against a rising sea level due to the Greenhouse Effect, people refuse to give up the coastal strip.

Of course there is no intention here of suggesting that all cultures in all times will share the same view of the hinterland. Obviously, there are numerous instances in which people will choose to live inland for direct access to fertile land, as well as for administrative, marketing, and traffic advantages, as explained by central-place theory (Boser 1968: 231ff).

The argument here is simply that, when people decide to live by the sea, they will be inclined to build as close to the water as possible. Beached here may do so because of land shortage. Fishermen may do so in order to minimize the distance from boat to house. There may be fish or mussel or turtle farms to serve, boats to build and watch, perhaps camp sites to watch, or merchant ships to be raided on the high seas. In some instances, particularly among affluent sub-cultures, people may live by the sea for aesthetic reasons – simply to feel the breeze, watch the sea, and have quick access to the water.

By combining this anthropological concept with the geomorphological beach profile, a specific area in the coastal zone is hereby established. This area is immediately behind the backshore, which extends from the upper limit of wave wash to the starting point of vegetation, a sand dune, or a sea cliff (Russett 1970: 12).

As a global model, the concept has three exceptions. First are salt marshes, which may be built along periodically flooded beaches, and therefore within the backshore. This is particularly a tidal-site or river phenomenon, such as the North-Isle lake sites in Switzerland and Italy (Russett 1972, 1988). A second exception are artificial islands which are still being built by the Cham Indians of Panama and the Marsh Arabs of Iraq. Archaeologically, the best known artificial islands are the *crannogs* of Scotland and Ireland. These consist of submerged boulder mounds kept in

place by underwater, again mostly a phenomenon of island nations (Merriman 1988). Finally, a third exception are coastal areas, where people may occupy rather than build, both on the nearby edge of the hinterland

Estimated Shoreline Migration

Estimated Shoreline Migration (ESM) equals the estimated distance that a given shoreline has either progressed or retreated, as measured from a known archaeological site (Figure 16). This is done by mapping the site's position within the area's beach profile. Based on the principle of coastal sediment patterns discussed above, ESM can be used as a fairly accurate measurement of shoreline migration. By correlating the location, age, and ESM of diverse archaeological sites, it is possible to reconstruct the relative sea level history of particular coasts and islands.

Eventually, once sufficient island sites are known—meaning “island” in the wider geological sense, which includes the coast as well as the submerged portions of continental and island shelves—it will be possible to identify the periods of polynesian both above and below sea level, and to predict the location of submerged land sites.

Lateral ESM, Spatial Information

Whether motion may be,
or whether still and again produced
also can become it?

The Bible, Ecclesiastes 1:10

Even upon the sea
motion is

and land's simply lateral

(Shakespeare/Teahupoo: 3.1.111, 140)

Marine transgression is a filter through which coastal sites may be modified but not necessarily destroyed. Contrary to the traditional view, there is growing evidence that local sites may survive transgression even in high energy areas. This seemingly paradoxical situation is due to the fact that most coasts, even high energy coasts, contain sheltered areas (Flannery 1983: 194). Seen through a LANDSAT photograph or small scale map (1:240,000+), the north shore of Puerto Rico appears as an exposed coast, which is certainly it is an overall sense. Yet, upon closer look (1:10,000 or ground truth), it becomes evident that even the rugged north coast has protected or semi-protected areas behind headlands, coves, or reefs. These sheltered spots were commonly favored by prehistoric peoples (Mearns and Flannery 1983).

Coastal Sites and Their Environments

Based on the comparative analysis of the prehistoric sites presented in Chapter 4, a number of significant patterns were recognized. Twenty-seven out of a total of twenty-nine sites were located in protected or semi-protected coasts, including nine out of eleven sites on the north coast.

The two submerged sites of Isla Verde and Cayo Batiques are both located on the western side of sandy spits, off the north and west coast respectively. These spits are protected by either beachrock (Isla Verde) or calcareous (Cayo Batiques), affording a barrier to both waves and ocean currents. It was interesting and rather surprising to realize that all six sites on spits are located off the east or south coasts, an area where Puerto Rico's insular shelf exceeds 10 to 15 miles. Without exception, these sites

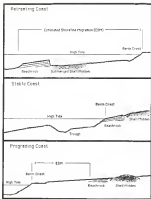


Figure 28. Geomorphological beach profiles of retreating, stable, and prograding coasts.

are separated from the main island by shallow water, and undoubtedly were tied to Puerto Rico at the time of prehistoric occupation.

Another surprising correlation has been noticed for the west coast. While the Colonial site of Jayade is subaerial, and the Colonial site of Caya Betances is submerged in shallow water, the contemporaneous site of Punta Ostionas is some 18 m inland. An evidence from the inland beachrock pavement, it is clear that sometime between prehistoric occupation and modern times, the coast has progressed at least 10 meters. Moving further south, the Colonial site of Boquerón is 40 m inland. On the south side of Boquerón Bay, the coast has formed an area of sedges or salt pits, as well as the large Boquerón Lagoon. Finally, at Caba Rojo, on the SW tip of the island, the present middle of Cayao is 1 km from shore.

Rouse (1944: 276) has written of Cayao: "Even there is no water supply in the vicinity, the Indians themselves must have relied on rainfall, unless they carried water from a distance." Rouse adds that the site is not particularly suitable for shell gathering, as it is "not directly on the shore, where one would expect it to be" (Rouse 1944: 281). Yet, shell was the staple and almost exclusive diet at Cayao, as no animal or fish bones have been found at the midden.

Rouse's interpretation is an excellent example of the traditional approach in Caribbean archaeology, which implicitly assumes a stable prehistoric topography. From a geoarchaeological perspective, it is clear that the southeast corner of Puerto Rico has been uplifted. This has resulted in significant changes to the topography as well as the hydrology, and presumably the cultural ecology of the Caba Rojo area. As the SW coast was uplifted, the streams that were cut within a short walking

distance of Cerros dried up, and the shoreline progressed, resulting in an island site.

The Cerros Controversy: A Short-handled Perspective

Brown's excavations at Cerros were the first evidence in favor of a prehistoric horizon in Puerto Rico. However, due to its location so far from shore, combined with its rather small size (20 m [66 ft.] in diameter), the archaeological community has been skeptical about a prehistoric "Cerros culture" (Davis 1946: 34, 35). Excavating at Cerros in the late 1930s, Brown carefully weighed the possibility of this being a camp of coastal Indians or even historic peoples. Yet, in spite of an informal report of isolated surface postholes, Brown found neither historic nor posthistoric ceramic artifacts at the site (Brown 1932: 375). In a cultural chronology of the Greater Antilles published in the early 1930s, Brown (1931: 351) included a prehistoric Cerros horizon for Puerto Rico, contemporaneous with the prehistoric Kona Bay Site in the Virgin Islands. A few years later, the prehistoric site of Cueva María de la Cruz was reported for Puerto Rico's north coast (Morgán et al. 1935). As previously described, this is a river site with a prehistoric level beneath an Igneri-Caracas contact level. Excavations at this site proved that Brown was correct about a prehistoric horizon in Puerto Rico. Yet, the Cerros site was virtually forgotten, excluded from Brown's later chronologies of the Greater Antilles (Crawson and Brown 1948, Brown 1944: 303, Fig. 5, Brown and Adams 1978: 404, Fig. 13.4). Following the non-controversial, prehistoric discoveries at María de la Cruz, Cerros's prehistoric west side story was relegated to a minor place.

But Bocas was right in the first place. Corozo is the median of prehistoric coastal dwellings; its antiquity is proven by 1 km of shoreline migration (Corozo ESM = +1 km). To illustrate the correlation between the antiquity of Corozo and shoreline migration, consider the nearest site of Bequeria, where the coast has progressed some fifty meters (Bequeria ESM = +50 m). In comparison, shoreline migration at Corozo depicts a 2250% increase over Bequeria, far too much to be explained by synchronous differences in silt and littoral transport. Clearly, there is a difference of time element involved here. Bequeria and Corozo belong to different time periods of Caribbean prehistory.

El Carrillo-Contador: A Geomorphological Perspective

As an additional example of the model's usefulness, let us consider briefly the later site of El Carrillo, in Caba Baja. This site was not included in Chapter 4 as it is a workshop rather than a habitation site, and therefore the principle of ESM cannot be applied as precisely as to shell middens. Approximately 1 km from the west coast, El Carrillo has been regarded as a prehistoric workshop roughly contemporary with the Marden-Barrera complex of the Dominican Republic, and tentatively dated at 4000 B.P. (Ortiz 1976, Pika and Ponsol 1974). This has caused controversy, as other researchers suggest that El Carrillo may well have been the workshop of late prehistoric) ceramic cultures (Davis 1968: 41). The present model supports Ortiz (1976, Pika and Ponsol (1974) may be correct in interpreting El Carrillo a prehistoric workshop, which was originally located closer to the 2000-4000 B.P. paleoshore.

On the north coast of St. Thomas, which apparently has undergone less uplift than SW Puerto Rico in posthistory, Loveland Quarry is located roughly 500 m (1,675 ft) inland from the presumed shoreline at Magen's Bay (Folios 1908: 540, Fig. 1). In terms of post-history, it makes perfect sense to imagine the people of Magen's Bay living by the seashore, and working in the nearby workshop of Loveland as needed. If this interpretation is correct, El Carrillo is simply a much more uplifted version of Magen's Bay, and there should be pronounced midline similarities between this workshop and the west coast. Assuming that beach ridges may be identified from aerial photographs, the "last midline" should be found on the inland side of these ridges.

Mid-Continents of Submerged Seas: Isla Verde Case Study

Of the two submerged seas of Isla Verde and Cape Raimon, the former has been extensively studied by the author (Wags 1966, 1982 and subsequent observations). Isla Verde is located off a coast currently undergoing severe erosion, and its work presents an interesting case study on submerged site modification.

The erosion rate of Punta el Medio has been estimated at 1.8 m/yr (Donaghy et al. 1998: 81). During a 300-day monitoring period (20 Jan. 88 to 27 Nov. 88), these researchers observed that erosion dominated from January to mid May, followed by accretion until late June, and then again accretion with some limited deposition from August to November.

As previously described, Isla Verde Sea is located on the eastern side of a large beachrock pavement. The midline is located 140 m (525 ft.) off the shore, at the top of Punta el Medio, at a depth of 4.2 m (13.8 ft.). The

beachrock pavement flanks the midden nearly on three sides, offering protection from both surf and the prevalent west-board longshore currents.

Without question, the surficial levels of the midden have been disturbed by sediment movements. This may be true of the entire site although the presence of *Elephas* postholes on the surface of the sea-berm, but not below, suggests otherwise (Wags 1981: 44). Either way, it is clear that this submerged site has survived transgression at least as an assemblage. Culturally-relevant information from this site can and has been extracted, very much as on a land site.

In principle, site modification factors at Isle Verde include 1) waves, 2) longshore currents, 3) marine animals, and 4) human activity. As previously mentioned, the site is located behind a small inlet, within a limited embayment partly blocked from the open sea by offshore reefs of volcanic origin. It is here, on the offshore reef, that the largest waves break. A smaller, second line of breakers rolls at the edge of the beachrock pavement, over 100 m (328 ft.) north of the site, as may be seen in aerial photographs (Wags 1981: 43, Fig. 10). Finally, small waves <0.5 m (2-6 ft.) break against the beachrock skirt immediately north of the site. In short, the site suffers little attack from waves.

The south or shoreward end of the midden may have drifted in a general westward direction, due to the effect of longshore currents. Partial erosion of the midden is also possible. Aerial photographs (1957 onwards), together with the author's personal observations (1972 onwards), demonstrate that the channel between the site and the shore has both widened and deepened through time. As an addendum to the 1970s, during low tide, the author and various friends once managed to wade across 50%

of the distance to the inlet near the site. Today, anyone under 2.4 m (7' 8.6") in height would not be able to repeat this "task."

Yet, the area immediately behind the midden has suffered relatively little modification other than the seasonal variation of the mangrove bank at the SE side of the inlet. In some ten years of careful observation, the author has noticed no significant changes in the midden's composition or physical appearance. Apparently, both the beachrock pavement and the adjacent reef act as a barrier to the adverse effect of ocean currents.

Recently, the passing of Hurricane Hugo offered a unique opportunity to study the site following extremely adverse conditions. While the eye of the hurricane struck the east coast, winds of 60 knots (mph) and up blew against the site area. On the east side of the embayment, at Isla Verde Beach, the palm and palm trees suffered such damage that the beach quickly reclaimed the place. It was as if an angry god had stolen the beach and replaced it with another. Where the trees should have been, there was only a wasteland as barren that it revealed the full landing strip of the airport behind it. At Punta el Miedo, we ran into a shorefront wall and scraped out parts of of a hotel terrace at the point. The grass at the margin of the inlet was literally uprooted by waves lashing against the bulk up coast. After the hurricane, the author snorkeled to the site with slabs on hand. The midden was intact. In the shallowest places, the seagrass had been cut down to the roots, but these roots were still there.

The report of mounds from an submerged land near a lagoon known At Isla Verde, small gobies (*Dobsonia*) and flatfish (*Stenodermatus*) partly bury themselves in mud over the site. As amazing as it may sound to non-divers, the author has seen an octopus collecting pottery shards at the site. As these various fragments are lighter than most stones of similar

properties, one possible hypothesis is that the natives preferred the lightweight potshards as housing material. Then again, perhaps he simply liked their texture. Fortunately, this unexpected type of potworking is restricted to surface collecting!

Worms are common (Eisenhammeria) may be seen with very small potshards attached to their spines, apparently as part of their normal life. *Ecobius* spp. (*Opisthotropis*), the false 'ear worms' of the Caribbean, often bury themselves in mud, with heads protruding out of the sea-floor. Deeper and generally permanent burrows are dug by Caribbean garden ants (*Myrmecochloa* spp.), although the author has not seen them at the site. Finally, there are ribbon worms, which may be found buried in muddy sediments (Bastall 1984 III, 22, 246; Carson 1982). Overall, the impact of these relatively small animals on the midden appears to be minimal if not nearly undetectable. Without question, these faunal disturbances more than likely compare with the impact on terrestrial sites of cattle, agriculture, mosquitoes, and land crabs, not to mention construction.

As for anthropogenic disturbances, at Isla Verde there are limited to sport divers. In more than twelve years of personal observations, the author has seen about an equal number of divers sharing in a massive ribbon turtle glaucon house, or perhaps a potshard, strigoid and jet diver ignoring not that they are swimming over the ruins of a native village. As the nature of these sites is better known, looting becomes a possibility. But that is a risk that must be taken. The looting of archaeological sites is a negative force to be fought with education and policing, not back-bash. The frequent situation of 'Juan's site,' which only Juan has seen and dug, is detrimental to archaeology and must be avoided.

Submarinal Sea Production

Now you are delivered by the sea on the depth of the waters your waves and all your company have given to us, make you. All who live on the mountains are supplied by you.

The Bible, Isaiah 27: 34

The sea is everything. It covers most lengths of the globe. It has both a face and body. The world is to speak, begin with the sea, and she knows but that it will also end at the end. There has happened something.

John Vance (1974: 18)

The comparative approach pursued in the previous section has opened a geomorphological vista that is all but invisible from the traditional sea level of analysis. Through the concept of *SSM*, correlations between sea level and prehistoric settlement patterns may now be explored within a strong empirical foundation. Clearly, the sea and the coast are a lot more than background scenery for Caribbean prehistoric excavation. Indeed, they are dynamic forces in constant change.

The Puerto Rico (Rapa Nui) Sea

The coast of Puerto Rico may be divided by shoreline types (Gardner and Glass 1967), by drainage basins (Morelock 1979:18), by beach systems (Morelock and Trumbull 1982: 147), by form, 'into five types of shoreline as six separate stretches' (Kaye 1980b: 51), by geographic regions (Pati 1974: 887), or by historic setting (Morelock 1979: 5), etc. For the geomorphologist, the usefulness of these geographic or geomorphic divisions depends on the questions being asked, as well as on the scale of the investigation (Hansen 1982: 43).

From the perspective of sea levels, tectonic history is by far the most relevant approach. Directly or indirectly, most of the environmental characteristics of the coastal zone are highly influenced, and in some defined, by tectonism. These characteristics include sedimentary and subsurface relief, as well as hydrology, soils, vegetation and aspects of climate (Pock 1974: 25).

In this section we will attempt to correlate the total set of archaeological sites and their ESM values, within a general consideration of the tectonic history of Puerto Rico. Essentially, this is similar to what was previously done for the YW coast, except that now the scale is larger in both chrono- and spatial terms. At this level of analysis, it is beneficial to present a wider geographic perspective, including Puerto Rico's offshore islands of Vieques and Culebra, as well as the nearby Virgin Islands. With the exception of St. Croix, these small islands are an extension of the Puerto Rico shelf, reaching as far east as Anegada Island (Dunn 1982).

Much of Puerto Rico, 2550 m (11,480 ft.) of tectonic subsidence has occurred since the Miocene (25-11 Ma B.P.), along the steep break of the Puerto Rico Trench (Parker et al. 1989: 152). The result has been a pelagic environment and a narrow north shelf, 2-4 km wide (MacLeod and Trumbull 1985: 187). As a working hypothesis, it is suggested that the tilting of the Puerto Rico Virgin Islands Shelf was by late proterozoic time. This explains: 1) the inland position of proterozoic gneiss on the north coast, such as Jales, Papayo, and Camuy, 2) the rare absence of proterozoic sites on the north coast, and 3) the relatively similar ESM values of Cretaceous marine sites around the island.

It is no coincidence that Morón de la Cruz, the only proterozoic site so far reported for the north coast, is located near the prograded scarp of the

Luxa River. This late prehistoric site has been radiocarbon dated at 1020 B.P. (Stearns and Allens 1978: 448). Stearns states that since sedimentation has kept up with the gradually rising sea level, as indicated by this site as well as by numerous beach bench ridges near the outlet of the Luxa River. These beach ridges, the most extensive on the island, increase in height and presumably decrease in age as they approach the present shoreline (Kaye 1994b: 118). Based on the position of Marín de la Cruz and the nearby town of Luxa Aldes (founded in the 18th century), Kaye has dated the oldest beach ridges at 3840 to 4870 B.P. Following the evolution from marsh to ever higher beach ridges, Kaye estimated that the oldest beach ridges were formed when the sea level was 1.8 to 3.4 m (6 to 11 ft.) beneath the 1950-2004. (Kaye 1994b: 118).

On the south coast, the prehistoric site of Cape Colwell is located on a small islet in a mangrove lagoon, and is currently undergoing marine transgression. This site has been radiocarbon dated at 2124-2174 B.P., and contains impressively made groundstone tools (Polun et al. 1975: 88). As described in Chapter 4, the island sites of Jilón, Papayan, and Corón contain no groundstone tools, further suggesting the considerable antiquity of these sites. Nevertheless, there may also be a tectonic differential behind the BSM values of Cape Colwell and the three island sites. As shown in Figure 9, a fault line cuts diagonally across the Bay of Jilón and past the shelf edge into the Caribbean Sea. This is one of three known faults that cut across the north shelf, undoubtedly introducing additional tectonic diversity within Puerto Rico's north coast.

The archaeological record of Vieques and St. Thomas indicates some tectonic uplift within the span of prehistoric and possibly ceramic cultures. For instance, the prehistoric midden at Cero Honda on Vieques's south coast

is roughly 76 m (250 ft.) from shore (Figueroa 1976, Brown 1983: 358). In October's south coast, the average Holocene site of Playa Camacho is some 35 m (115 ft.) inland. On the north coast of St. Thomas, the cluster of prehistoric sites at Mager's Bay is some 55 m (184 ft.) from shore (Tidson 1976: 344, Fig. 1). Also on St. Thomas north coast, the prehistoric sites of Krum Bay, Guanchokela Hill, and Canelo Hill are roughly 12 m (33 ft.) from shore (Gross 1978, Ramsey 1940: 154). It is interesting to notice that the average moderns of Hall Bay, near Mager's Bay, are closer to the intertidal zone than the latter, suggesting a prograding coast up to the present age. In other words, the cultural sequence is reversed, with the older prehistoric moderns further inland than the modern ones.

In Vieques, an intensive survey prepared for the U.S. Navy (Friedman et al. 1984a, Part II, Vol. 1: 6-17) indicates that most coastal sites are presently located over 180 m (590 ft.) from shore, further supporting the hypothesis of uplift east of Puerto Rico. This is not to say that all Vieques coastal sites are inland, as at least two oceanic sites—San Chas and Parameyito 3—are presently within the intertidal zone (Friedman et al. 1984b: Volume II).

At this point, the landlocked archaeologist could surmise that there are no further prehistoric sites to be found in the Virgin Islands, as both coastal and prehistoric sites have been located inland in Vieques and St. Thomas. That, however, would be a bad case of archaeological myopia. We must complete our geoarchaeological tour of the Puerto Rico-Vieques shelf before plotting the archaeological data within a sea level framework.

Off Vieques Goods, NE of St. Thomas, at least one prehistoric submerged site has been found by deep-sea operator treasure hunter Bart Kildea, yielding "darkly polished Stone Age implements" (Thrupp 1988

1840). As these are groundwater artifacts, most likely petiolated coifs, it may be tentatively assumed that the discovery is questionable either a constant or low groundwater mound. Further north, near the NE tip of the Puerto Rico-Virgin Islands Shelf, lies Anapoda Island. Being about 6.5 m (21 ft.) above sea level, Anapoda is so low that at times it becomes almost invisible to Atlantic sailors approaching the Caribbean in rough weather. Hundreds of ships have come to a wreck here and on the northern fringes of Anapoda Reef. Some sailors call this island "The Low Land." And some, perhaps not realizing the wisdom of their words, call her "The Sinking Land." In Spanish, *Anapoda* means "The Flooded Island."

Figure 15 presents a preliminary geoarchaeological plotting of tectonic movements for Puerto Rico and the Virgin Islands. The map is based on the ERM values of the archaeological sites described above, as well as on some geomorphological considerations. The preliminary nature of the map must be stressed, as it is intended primarily as a heuristic tool to illustrate the regional, super-site application of ERM.

The tectonic tilt of the Puerto Rico-Virgin Islands Shelf is clearly reflected in the archaeological record. By comparing the age, locusts and ERM values of archaeological sites, we can define the tectonic and surface parameters necessary for any realistic reinteractions of the sea-level history.

Without exception, the petrousites discovered in Puerto Rico, Vieques and St. Thomas are relatively recent, all of them dated at or under 5,000 B.P. (Dunn and Adams 1979: 644, Table 13.4). For older sites, we can only assume negative evidence. In Puerto Rico, early petrousites sites compatible to Haysenaki's Mardita, Camorra and Osborn have not been found. As previously stated, parts of Haysenaki's north coast have been

uplifted well into the Pleistocene Epoch, as evidenced by raised coral reef terraces (Alexander 1955: 163). The opposite is true for the north coast of Puerto Rico, which is a geologically unstable coast. If early postcolonial peoples occupied the north coast of Puerto Rico, structures are certainly not to be found inland. Undersea sites are the only alternatives. As Shalicki Holmes would put it: "When you have excluded the impossible, whatever remains, however improbable, must be the truth."

Data from historic sites and documents indicate that the sea level has risen at least 75 cm (25 ft.) since the early 19th century. Shoreline migration at Aguadilla Fort San Francisco is clearly the result of sea-level movements precisely dated to 1816. However, the other historic sites on the north coast indicate a rise, but their exact rise in sea level. Otherwise, if the slight submergence of the north coast in historic times were uniformly isotonic, we would expect to see greater variations in the DSM values of prehistoric ceramic sites throughout the island. This is not to deny testimony in historic times, but simply that the relative level of the sea has risen in the last 200 years. The constant change is there, but it is there.

Suriken Land Sites: A Caribbean Model

A model of Caribbean submerged land sites is presented in Figure 19. The model is self-explanatory, but a few comments are in order. Shalicki sites were estimated from the dimensions of known prehistoric sites, including Cane Head (28 X 8 m), Canees (29 X 22 m), Cays Cofres (30 X 20 m), Magu's Bay (23 X 9 m), and Jolón (15 X 25 m). These last two sites come in diameters of 4 and 5 madden, respectively, suggesting that larger



Figure 17 Preliminary map of submerged land sites for Puerto Rico and the Virgin Islands, based on section 145

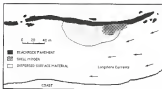


Figure 18 A model of prehistoric, submerged land sites for the Caribbean

larger mounds would have resulted in a sectorized deposit. For the model, a 10 X 10 m (33 X 33 ft.) mound was assumed.

Mound shape is assumed to be the combined result of jet placement and the configuration of the river. Sites located at head points such as Iste Yorda, Punta Manayon and Punta Corcos will tend to be circular or oval, while sites on straight segments of coast may produce elongated mounds. Of course, the shape of the mound may be transformed by post-depositional processes before or after marine incursions.

Artifact and profile preservation at submerged land sites should not be regarded as poor, as land-based archaeologists often assume. Stone, shell, pottery, and bone material have been found in good condition at Iste Yorda. Moreover, shards with the typical Gifford red slip have been retrieved both above and below the surface of the mounds. The only obvious evidence of artifact damage at Iste Yorda consisted of a few, low-fired, coarseware shards, which crumbled after removal from the sea water, presumably through salt crystallization (Sly 1943: 30). Comparable damage has been observed on coarseware shards excavated on land. Invariably, these are low fired wares.

Polish and phytolith profiles of submerged land sites may be retrieved through coring or during excavation, as on a dry-land site (Clifford 1943: 173, Rappé 1944a: 41). Submerged land sites may be partly sealed in anoxic conditions, particularly in mud or slough deposits. In these situations, virtually anything may be preserved for millennia, including wood, stone and bone. At the submerged site of Tybjerg Vig, off the Danish island of Fynen, wooden artifacts dated to 7000 B.P. have been recovered in good condition (Andersen 1948). Such preservation is very seldom found on land sites.

Conclusions

What was the shape and size and color and tone of that old's experience? We stopped here a few times and gave the thought of it, related as some while ago, to the past and future, related to the body, related to the living nature and the world.

John Barlowe (1970-1972)

My friend, My friend,
It was known the first year,
The first year's effort,
Out of the end of the world's end.
Faint, 1970-1972

The geospatiality of coastal change provides a new framework for Caribbean prehistory. The ramifications that may be perceived are vast, both above and below water. Evidently, early prehistoric aborigines did not carry their shellfish for thousands of miles inland, neither did they sit around waiting for a little rain to quench their prehistoric thirst. They simply adapted to new ecological conditions that were either created by themselves, such as over-exploitation, or created by strictly environmental, non-ecological variables such as hurricanes, tectonic changes in sea level, climatic changes, etc.

Through tectonic uplift, many streams on Puerto Rico would most probably dried up. When that happened, the people packed their possessions and moved to a new site. On the north coast, the situation was different (Figure 10). Here the sea rose a little more every year, through subsidence as well as through a rising sea level. Streams may have gradually altered their course, perhaps meandering, profusing marshlands and small lagoons in some areas, but never drying up.

Considering that the elevation of Carac is 5 m (16 ft.) above present sea level (Beane 1952: 279), the calculation of the north coast may be estimated at roughly 5 m. since late prehistoric times (3500-3580 B.P.). Adding a conservative estimate of 5.5 m (18 ft.) of eustatic rise in sea level since then, one might expect to find prehistoric sites off the N coast at an average depth of 10.5 m (35 ft.), as presented in Figure 20. Of course, Puerto Rico is not the perfectly rigid slab that is implied in these preliminary calculations. At this point, the conceptual framework is more important than the exact measurement of tectonic tilt.

For the inland areas of the north coast, we may estimate the ratio of sea level to elevation magnitude somewhere between 1:100 to 1:1000. Sticking again to the preliminary nature of these calculations, we arrive at an estimated elevation magnitude of 550 to 5480 m (1,800-18,000 ft.) for the north coast, near 5,000 B.P. Currently, that takes us to the offshore volcanic ridges by the 5.5 m (18 ft.) estimate. It is worth noting that this possible correlation may be no more forced, but rather come out as a small corroboration, as the estimation given above was based strictly on the archaeological data.

Estimated Statistics

Underwater archaeology need not be expensive. Serious research can and has been accomplished on minimal budgets, as demonstrated by Fleming (1961) and Gifford (1962) in the Aegean Sea, by Eklund (1961, 1968) off the Mediterranean coast of Israel, by Weisbach (1964) off Colombia; by Boyet (1960a) off West Florida; by Maurer (1962) off Southern California, etc.

The authors own underwater excavations at Isle Verde was funded through a small donation, volunteer work, and pocket money. Excluding the Varques Cubates Pipeline Marine Survey, which involved a magnetometer, underwater video, a fairly large crew, etc., the dive operations behind this study were extremely simple and inexpensive. In many cases, equipment consisted of overbelling gear, a slate and a segmented rod.

In the United States, as required by the National Historic Preservation Act of 1980, and the Outer Continental Shelf Lands Act Amendments of 1974, leasing for offshore oil and gas exploration must be preceded by a competent, marine-archaeological survey. Ironically, not one submerged land site has been located through these costly OCS surveys (Masters and Fleming 1982: 806). This is not to say that the OCS surveys are a waste of money. In addition to possible shipwreck sites, these surveys provide submarine geophysical data that will be of use to marine geoarchaeologists in the near future. In the author's case, the Varques Cubates Survey allowed the exploration of a 1.5 m trench (10 meters wide) of sea bed across Varques Beach. As previously mentioned, this resulted in the discovery of a headstock pavement at a depth exceeding 30 m (30 ft), 4 km away from the nearest island.

It should be possible to detect densely-packed middens through sonar, speedily with sub-bottom profilers. An experiment could be set up by reading the 'signatures' of known submerged middens and then testing for other sites in high probability areas, such as Ponta Rasa north shore, Farolão Tromps Bay (Warren 1980), etc. At present, however, the best approach to finding submerged land sites also happens to be the disquieting and rather drying

The Caribbean is one of the most popular dive destinations in the world. Virtually all the islands have diving enthusiasts. Thousands of divers on vacation from the U.S. island visit the region on a yearly basis. As a scuba instructor, the author is quite aware of the untapped resources that sport divers represent for archaeology. Given a certain amount of "adventure" or "exotism," divers are more than willing to volunteer their leisure time to archaeological research. In the long run, divers will either become a positive or a negative force regarding archaeology. It is essential for archaeologists to convey the importance of submerged sites to divers, and to incorporate interested divers into their work.

Archaeologists in search of submerged land sites should keep their eyes and ears open to anything related to the sea and the coast, sand formations, offshore dredging operations, biological studies, dive shop rumors, cruise studies, shipwreck archaeology, treasure hunting, coastal and offshore construction, etc. Typifying may also be a useful lead. Any mention of points, keys, or reefs related to the word "shoalish" may indicate the presence of an intertidal or partly submerged midden. In some cases, small keys may well be the remains of densely packed mounds with a higher resistance to erosion than the surrounding terrain. In sandy areas, patches of sea grasses often indicate the presence of rocks, wreck matter or perhaps a submerged midden beneath a sand veneer. Of course, there are areas where sites may be buried under tons of sediments, in which case they may be never be found, except through the chance hit of trawling operations for (shoalish) or ointempral purposes.

Submerged beachrock and volcanic formations are clear indicators of paleoshores. The inland side of these formations is a good place to check for middens, particularly in the vicinity of submerged river banks. Off West

Florida, Ruggel (1946a: 26) has followed the random course of a river for over 1 km (1/2 natural mile) into the Gulf of Mexico. His technique – a relatively expensive, according to his assistant

The importance of inland sites must not be discarded. Although the coastline is the primary place to search for the Caribbean's early prehistoric sites, upland areas provide a setting in which to find sites dating as early as 3000 B.P. in Puerto Rico, and possibly 4000 B.P. in Hispaniola (Ortega and Gussner 1961, Veloz and Ortega 1976). Inland shell middens – particularly prehistoric ones – which are away from rivers large enough for canoe navigation, should be regarded with attention to possible uplift. Of course, rivers may change their courses or dry out due to these same tectonic movements, or to new climatic regimes, but any stream large enough for canoeing will leave its imprint on the ground.

For Puerto Rico's south coast, it is fundamental to explore to what extent uplift kept up with the rising sea level. Presumably, coastal submergence delayed the south coast's tectonic uplift until 7000 B.P., and perhaps up to 5000 B.P. This is presented as a working hypothesis, or rather a strategy, in order not to lose sight of the possibility of early prehistoric sites submerged off the south coast. Certainly, an alternative hypothesis of early prehistoric inland sites should be kept in mind for Puerto Rico's upland south coast.

San Juan Island, Antigua

Based on the results of the present study, and on the preliminary interpretation of published archaeological and geomorphological data from other parts of the Caribbean, the overall tectonic trend appears to be a NE

10). This is in agreement with Caribbean plate motion estimates derived from the slip vectors of shallow earthquakes, which indicate an ENE plate movement of 4 cm/yr for the past 7 Ma. (Gyllen et al. 1992: 10,000). As a preliminary hypothesis, it is suggested that, for the Lesser Antilles island arc may be found on ENE trends, with the greatest possibility of submergence west of the SE coast. West of Puerto Rico, the tectonic setting is far more complex, with Cuba outside of the Caribbean plate proper, and sea floor spreading along the Cayman Trough. Uplift is evident on the north coast of Grand Cayman (Emery 1955) and Jamaica (Jones 1965) while Cuba and Hispaniola show a highly complex history of folding, faulting, depression, and uplift (Alexander 1953; Reid 1986a).

The possibility of Caribbean, submergent processes is evident, even from tectonic history alone. In spite of the outstanding number of archaeologists who have worked in the region, we still have a very superficial notion as to how the archipelago looked three, five, seven, ten, twelve, or fifteen thousand years ago. A complete model of coastal change, such as the one presented in Chapter 5, can only be regarded as a first step. There is an archaeological need to reconstruct the tectonic setting of the entire archipelago.

The concept of ESM offers a new approach to the dating of archaeological sites. Intrinsically separate from radiometric dating and cultural chronology, ESM allows us to perceive spatial-temporal relationships beyond the scope of the other dating methods. In other words, the dating of a site becomes inseparable from the larger ecological framework of environmental change. Just as, for instance, ceramic typology provides a culturally-oriented dating method, ESM provides an environmentally oriented dating method.

A changing sea level may imply vast environmental transformations. Reduced coastal plains may trigger population pressures (Barfield 1988), the depth, temperature, and salinity of lagoons and estuaries may increase or decrease beyond the tolerance of numerous animal species (Longheart et al. 1987, Uhlir 1997, Walsh 1998), coastal lowlands may be transformed into swamps, which may in turn become breeding grounds for the malaria-carrying, anopheline mosquitoes, changes in elevation may produce alterations in temperature and moisture regimes, indirectly affecting the flora and fauna (Lyons 1975: 460).

What was the impact of tectonic and eustatic movements on the topography and ecology of Caribbean islands? And what role did these environmental transformations play in prehistoric change? We must answer these questions not only for the coastal regions that we see today but also for those lands that are now beneath the sea. The Caribbean possesses a rich, untapped, archaeological landscape that must be explored. As you men Gato Barbieri once said of man: It has limits, but we do not know them.

Caribbean Maritime Hunter-Gathering

The present study suggests that prehistoric hunter-gatherers ventured from the continental land masses in search of large marine fauna to hunt, and of new coastal environments for gathering. There is no question that the prehistoric shorelines changed numerous times over time. We also know that at least one stranded whale was consumed by prehistoric hunters in Hispaniola (Vries and Orlage 1978: 154). But is this an isolated historic event, or evidence of a tropical, maritime, cultural specialization process?

Wing (1979: 22) suggests that polynesian maritime lifeways require "both luck and skill to avoid the unpredictable elements: wind, and waves that may destroy boats and fishing equipment." This writer would agree; if the words "luck" and "unpredictable" were replaced with "knowledge" and "largely predictable."

The sea will always possess an element of surprise. The best diver may be attacked by an old or angry shark; the best captain may lose both boat and life to a hurricane or waterspout. The author can attest to this through both personal experience and diligently endured research. But there are extraordinary events that may occur once in a lifetime.

Most sea conditions are highly predictable, if people only know what to look for (Bellwood 1979: 297, Lewis 1972, Thomas 1982). The color of the sea and the clouds, the direction, temperature and strength of the wind, the movements of birds, fish, and porpoises, the patterns of sea weed on the surface of the water, all of these speak a secret language to be discovered by the careful observer. If one can learn the basics of sea prediction in a few years of observation, consider how much could be learned in centuries of accumulated cultural experience, and this not simply for the beauty of it and the quest for knowledge, but for sheer survival. If the secret loss of Caribbean polynesian navigation is now unrecognized by anthropology, it is not because such loss was poor or nonexistent, but rather because the evidence that existed is did not survive to tell us about it.

The hypothesis of sea hunters does not preclude the hunting of land fauna such as sloths, tortoises, rodents, etc. There is no cultural law that forces people to choose between the sea and the land. In the Greater Antilles, a number of bands may have chosen to live in the hinterland, perhaps with little or no contact with the coast. However, if such an Antillean "Pala-

Indian' because over time it was probably short lived, for the island's big game was virtually restricted to a few species of sloths (Olson 1973: 1952). In the prehistoric Caribbees, most big game was to be found in swamps, in wetlands, and at sea. If there was ever a big game hunting tradition on the Caribbean Anthropology, it was almost certainly related to the sea.

Caribbean archaeologists need to develop methods specifically designed for the region's tropical, maritime environment. The present study has explained the historical and ecological parameters of Caribbean sea levels. A methodology for reconstructing paleoenvironments has been presented. The next logical step is to search for submerged prehistoric sites off the north coast of Puerto Rico, and to take a closer look at the prehistoric, island sites on the south. Sites like Jolón and Dorcas should be "rescued" and dated through radiocarbon testing. The islands from Vieques to Anegada are the natural extension of the present study, as is, in a wider sense, the entire endogeography. By rethinking Caribbean prehistory in terms of plate tectonics and sea level, a fuller, more dynamic view of tropical ecological history will begin to emerge. Within this framework, it should be possible to progressively assess relationships between the prehistoric environment and the people who lived off it.

Sea Level Chronology: Practical Applications

Archaeological methods have a wide range of uses that we are only beginning to appreciate and explore. It is clear that archaeology can play a key role in reconstructing the sea level and historic history of the Caribbean and elsewhere. Certainly the time scales of archaeology and geology are different, but 7000 or more years of prehistoric human occupation is an

ample margin for detecting and measuring Holocene wetness and various trends. This is a potentially symbiotic relationship: for geomorphologists and geomorphologists may have a lot to say to archaeologists both in terms of methods and perspectives.

Coastal geomorphology can be highly relevant to urban planners, particularly now that we debate the possible, global consequences of the Greenhouse Effect. A by-product of industrialism, the Greenhouse Effect is caused by the release of carbon dioxide and other gases into the atmosphere. Like the glass enclosure of a greenhouse, these gases retain heat, therefore increasing atmospheric temperatures, and possibly triggering a glacio-isostatic rise in sea level. As in the reconstruction of past sea levels, geologists have not reached a consensus on the pace and potential threat of the Greenhouse Effect. This is partly due to difficulties in accurately measuring sea level changes, and in separating natural from tectonic factors. However, there is substantial evidence that the sea is indeed rising on a global scale. In Bangladesh, which is also extending due to tectonic depression at the mouth of the Ganges River, some researchers fear that, if the current trend in atmospheric warming continues, as much as 10% of that country's land will be lost to the sea in the next century (Rahman 1990). Considering that Bangladesh is already overpopulated, and that a rising sea level may also impact the groundwater with saline intrusion, it is clear that we are dealing with a potential problem of enormous proportions.

At the other end of the socioeconomic spectrum, a rising sea level implies the eventual destruction of coastal properties valued in billions of dollars. More importantly, the rising water may trigger ecological disaster due to the increasing number of toxic waste sites in the coastal zone. In

Holland, the sea must be sought in order to save the most productive farmlands in the country. In places like Louisiana and Texas, a shoreline retreat implies the loss of considerable offshore oil rights, measured mostly in three miles from shore (Kushlan and Pilley 1999). In short, sea level and shoreline migration go far beyond the academic problems of prehistorians, geographers and geomorphologists. Directly or indirectly, sea level is as crucial to modern humans as it was to the first prehistoric navigators who paddled their canoes towards the grey profile of a distant island.

Archaeology can and should play a role in the study of sea levels. Our targets are the countless prehistoric and historic sites that dot the world's littoral zones. These sites have a culture, as well as an environmental history to tell. If we don't face this challenge, neither discipline will borrow our methods and do it for us. To archaeologists, sea level research implies three potential benefits: I) making a contribution to the understanding of a practical, global problem, II) tapping a new source of funding for archaeological research, and III) working side by side with earth scientists. In the end, to paraphrase John Steward, it is all one thing, one planetary history, one fundamental reality. The prehistoric hunter-gatherers of tropical islands, the monumental destruction of the Spanish Moss, the millionaires in their remote villas, wondering if the state will pay for seawall repairs, the Marsh Arabs of Iraq in their artificial islands—they are chapters of the same human journey. We may think of these lifeways as belonging to different paradiises, different worlds, but the sea knows better.

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BIOGRAPHICAL SKETCH

Josua Vega is a marine archaeologist, archival researcher, professional diver, and writer who has participated in numerous shipwreck projects in Florida, North Carolina, Virginia, Belize, Spain, and Puerto Rico. He has over seven years of experience in Spanish paleogeography, including three years in the archives of Seville, Madrid, Salamanca, and Barcelona. In 1988, he directed the excavation of a submerged prehistoric basket off Puerto Rico's north shore. In 1984, he taught world prehistory at the University of Florida and participated in the excavation of an estimated 12,000-year-old ball site at Central Florida's Aucilla River. In 1994, he worked as a consultant in Spanish paleogeography at the P. K. Yonge Library of Florida History. Back in Puerto Rico, he directed the Vieques Sound Marine Archaeological Survey in 1997, ensuring that the construction of Culebra Island's water pipeline would not destroy any historic shipwrecks in the Sound. From 1985 to 1989, he directed the search for the 1558 shipwreck of the *Santa Maria de Jesus*, described in Spanish documents as "the richest ship ever to have sailed to the Indies" up until that time. After a year of preparations and four arduous months of marine work, employing sophisticated electronic instrumentation, he located the wreck in Puerto Rico's territorial waters. As a freelance writer, his credits include fiction in *Miami Water Magazine*, and four cover articles in the *San Juan Star Sunday Magazine*. He currently works as a bilingual copywriter at one of

the largest advertising agency in Puerto Rico, teaches special courses in archaeology, shamanism, and dangerous marine life at the Centro de la Nueva Educación, and is organizing the Aquarius III Foundation, whose goal is the fusion of art, ecology, archaeology, aesthetics, and underwater cinematography in the Caribbean. His other interests include surfing, hiking, tracking, rappelling, spelunking, chess, music, film, and martial arts. He holds a B.A. in anthropology from the University of Puerto Rico, an M.A. in anthropological archaeology from Florida Atlantic University, and is a certified scuba instructor with over twenty years of diving experience.

I certify that I have read this study and that in my opinion it conforms to acceptable standards of scholarly presentation and is fully adequate, in scope and quality, as a dissertation for the degree of Doctor of Philosophy


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